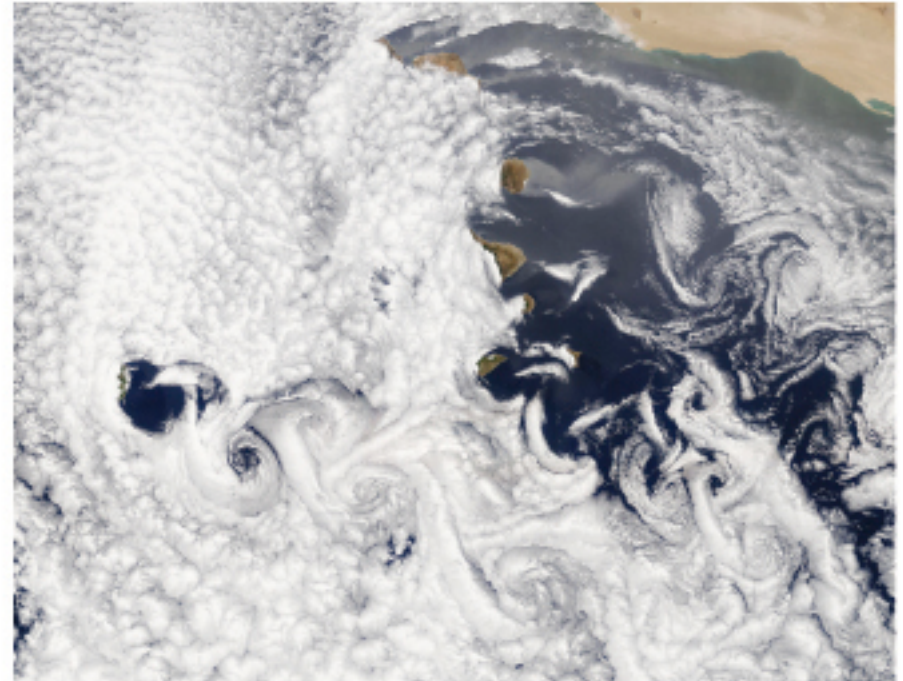


Ocean Modeling - EAS 8803

*Geostrophy and
Shallow-water model*

Introduction to
Geophysical Fluid Dynamics
Physical and Numerical Aspects



Benoit Cushman-Roisin and Jean-Marie Beckers

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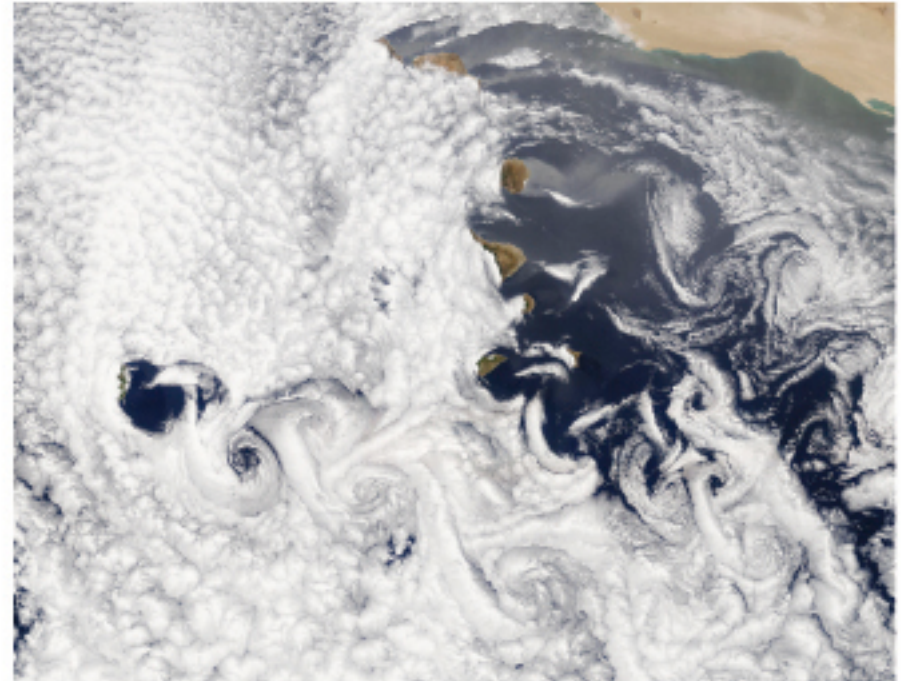
Chapter 7

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- Further scaling of the primitive equations of motion

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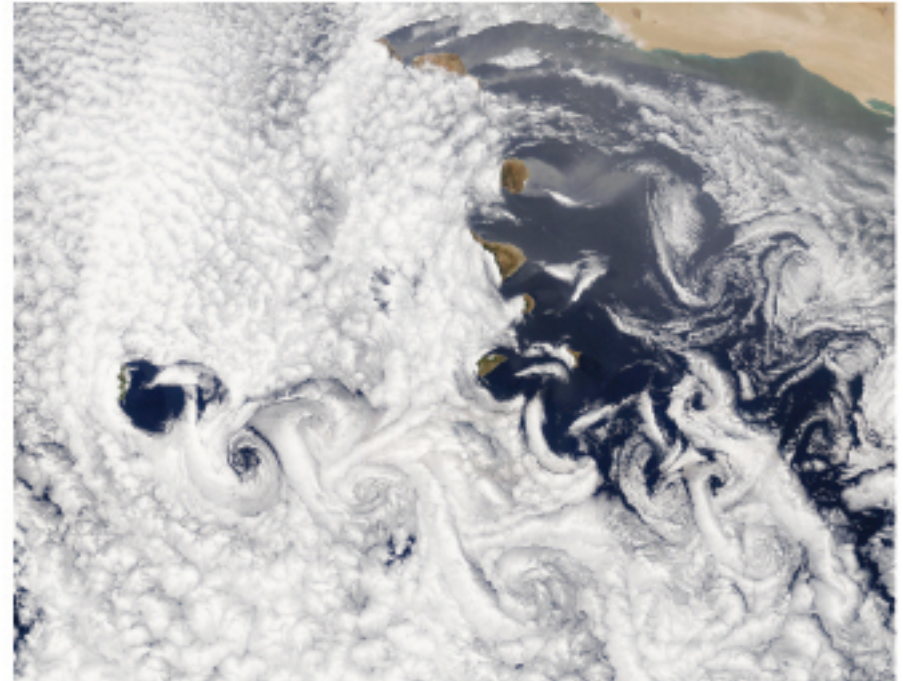
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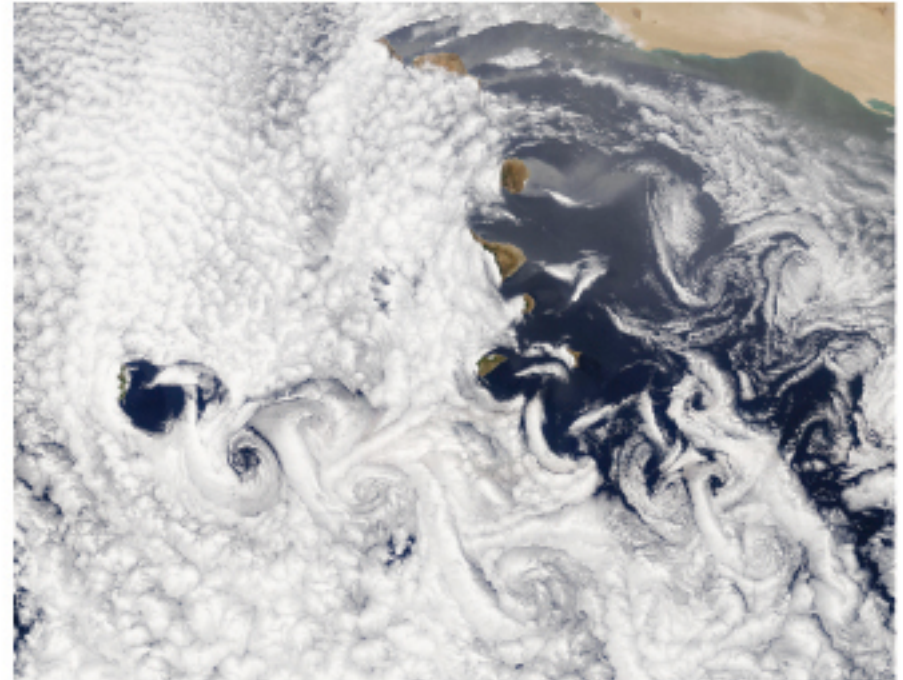
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Chapter 7

Geostrophy and Shallow-water model

- Further scaling of the primitive equations of motion
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- Steady and homogenous flows (*geostrophic flow*)

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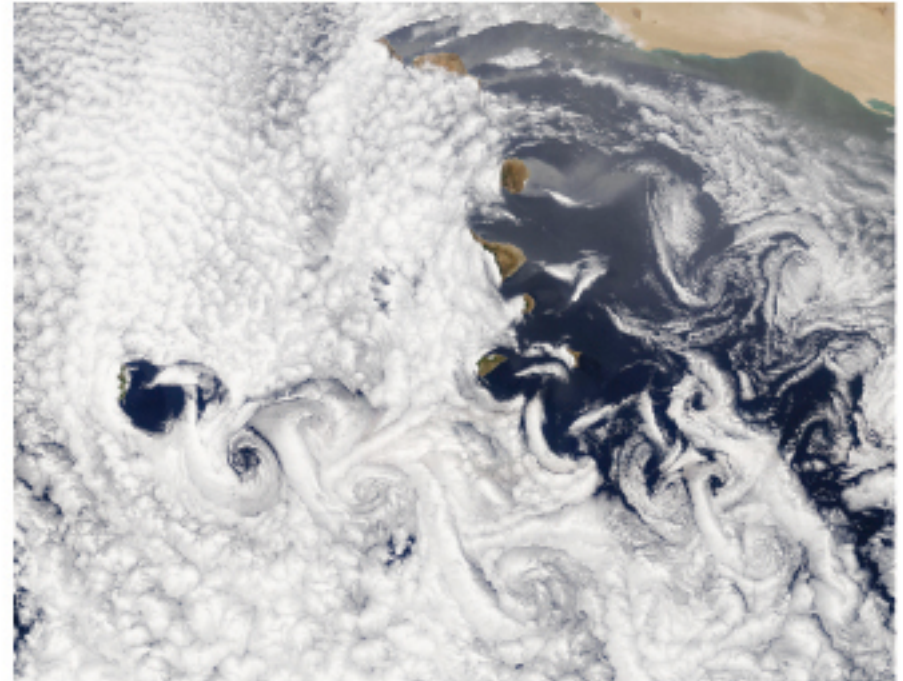
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Chapter 7

Geostrophy and Shallow-water model

- ➊ Further scaling of the primitive equations of motion
- ➋ Dimensionless number and active dynamics
- ➌ Steady and homogenous flows (*geostrophic flow*)
- ➍ Properties of geostrophic flows

Introduction to
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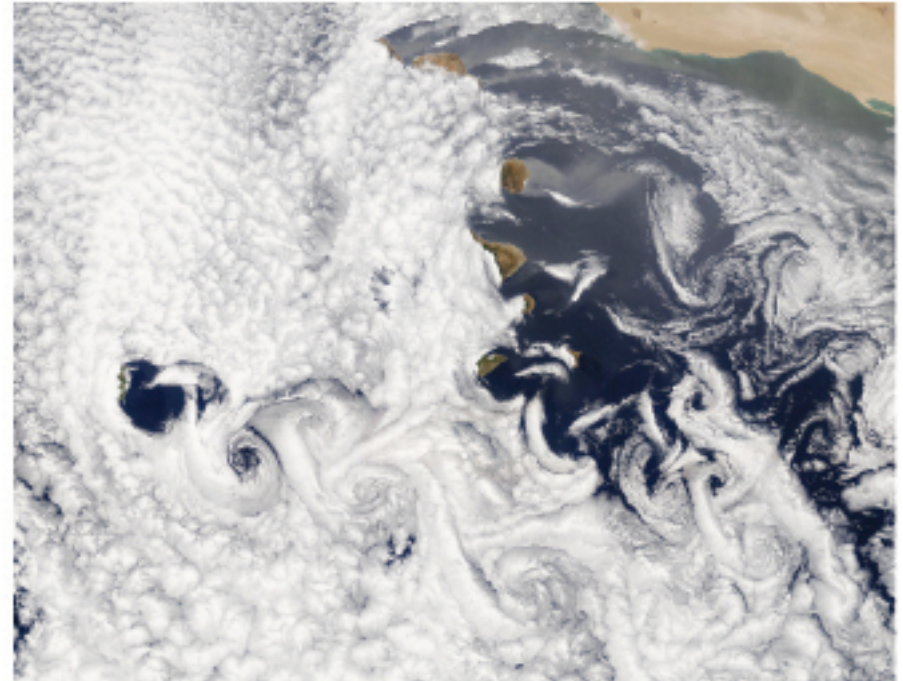
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Chapter 7

Geostrophy and Shallow-water model

- ➊ Further scaling of the primitive equations of motion
- ➋ Dimensionless number and active dynamics
- ➌ Steady and homogenous flows (*geostrophic flow*)
- ➍ Properties of geostrophic flows
- ➎ Unsteady and homogenous flows (*shallow water model*)

Introduction to
Geophysical Fluid Dynamics
Physical and Numerical Aspects



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Chapter 7

Primitive Equations describing motions of geophysical flows

x – momentum:
$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z} - fv =$$
$$-\frac{1}{\rho_0} \frac{\partial p}{\partial x} + \frac{\partial}{\partial x} \left(\mathcal{A} \frac{\partial u}{\partial x} \right) + \frac{\partial}{\partial y} \left(\mathcal{A} \frac{\partial u}{\partial y} \right) + \frac{\partial}{\partial z} \left(\nu_E \frac{\partial u}{\partial z} \right)$$

y – momentum:
$$\frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} + w \frac{\partial v}{\partial z} + fu =$$
$$-\frac{1}{\rho_0} \frac{\partial p}{\partial y} + \frac{\partial}{\partial x} \left(\mathcal{A} \frac{\partial v}{\partial x} \right) + \frac{\partial}{\partial y} \left(\mathcal{A} \frac{\partial v}{\partial y} \right) + \frac{\partial}{\partial z} \left(\nu_E \frac{\partial v}{\partial z} \right)$$

z – momentum:
$$0 = -\frac{\partial p}{\partial z} - \rho g$$

continuity:
$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0$$

energy:
$$\frac{\partial \rho}{\partial t} + u \frac{\partial \rho}{\partial x} + v \frac{\partial \rho}{\partial y} + w \frac{\partial \rho}{\partial z} =$$
$$\frac{\partial}{\partial x} \left(\mathcal{A} \frac{\partial \rho}{\partial x} \right) + \frac{\partial}{\partial y} \left(\mathcal{A} \frac{\partial \rho}{\partial y} \right) + \frac{\partial}{\partial z} \left(\kappa_E \frac{\partial \rho}{\partial z} \right),$$

Primitive Equations describing motions of geophysical flows

x – momentum:

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z} - fv =$$
$$-\frac{1}{\rho_0} \frac{\partial p}{\partial x} + \frac{\partial}{\partial x} \left(\mathcal{A} \frac{\partial u}{\partial x} \right) + \frac{\partial}{\partial y} \left(\mathcal{A} \frac{\partial u}{\partial y} \right) + \frac{\partial}{\partial z} \left(\nu_E \frac{\partial u}{\partial z} \right)$$

Primitive Equations describing motions of geophysical flows

x – momentum:

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Scaling of terms

$$\frac{U}{T}, \quad \frac{U^2}{L}, \quad \frac{U^2}{L}, \quad \frac{WU}{H}, \quad \Omega U, \quad \frac{P}{\rho_0 L}, \quad \frac{\mathcal{A}U}{L^2}, \quad \frac{\nu_E U}{H^2}$$

Primitive Equations describing motions of geophysical flows

x - momentum:

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z} - f v =$$
$$-\frac{1}{\rho_0} \frac{\partial p}{\partial x} + \frac{\partial}{\partial x} \left(\mathcal{A} \frac{\partial u}{\partial x} \right) + \frac{\partial}{\partial y} \left(\mathcal{A} \frac{\partial u}{\partial y} \right) + \frac{\partial}{\partial z} \left(\nu_E \frac{\partial u}{\partial z} \right)$$

Scaling of terms

$$\frac{U}{T}, \quad \frac{U^2}{L}, \quad \frac{U^2}{L}, \quad \frac{WU}{H}, \quad \Omega U, \quad \frac{P}{\rho_0 L}, \quad \frac{\mathcal{A}U}{L^2}, \quad \frac{\nu_E U}{H^2}$$

inertial terms

Primitive Equations describing motions of geophysical flows

x - momentum:

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z} - fv =$$
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Scaling of terms

$\frac{U}{T}$	$\frac{U^2}{L}$	$\frac{U^2}{L}$	$\frac{WU}{H}$	ΩU	$\frac{P}{\rho_0 L}$	$\frac{\mathcal{A}U}{L^2}$	$\frac{\nu_E U}{H^2}$
<i>inertial terms</i>				<i>rotation</i>			

Primitive Equations describing motions of geophysical flows

x – momentum:

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Scaling of terms

$\frac{U}{T}, \frac{U^2}{L}, \frac{U^2}{L}, \frac{WU}{H},$	$\Omega U,$	$\frac{P}{\rho_0 L},$	$\frac{\mathcal{A}U}{L^2}, \frac{\nu_E U}{H^2}$
<i>inertial terms</i>	<i>rotation</i>		<i>frictional forces</i>

Primitive Equations describing motions of geophysical flows

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Scaling of terms

$\frac{U}{T}, \frac{U^2}{L}, \frac{U^2}{L}, \frac{WU}{H},$	$\Omega U,$	$\frac{P}{\rho_0 L},$	$\frac{\mathcal{A}U}{L^2}, \frac{\nu_E U}{H^2}$
<i>inertial terms</i>	<i>rotation</i>		<i>frictional forces</i>

divide by ΩU

Primitive Equations describing motions of geophysical flows

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Scaling of terms

$\frac{U}{T}, \frac{U^2}{L}, \frac{U^2}{L}, \frac{WU}{H},$	$\Omega U,$	$\frac{P}{\rho_0 L},$	$\frac{\mathcal{A}U}{L^2}, \frac{\nu_E U}{H^2}$
<i>inertial terms</i>	<i>rotation</i>		<i>frictional forces</i>

divide by ΩU

$\frac{1}{\Omega T}, \frac{U}{\Omega L}, \frac{U}{\Omega L}, \frac{WL}{UH} \cdot \frac{U}{\Omega L},$	$1,$	$\frac{P}{\rho_0 \Omega L U},$	$\frac{\mathcal{A}}{\Omega L^2}, \frac{\nu_E}{\Omega H^2}$
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Important Dimensionless Numbers

measuring the sizes of the terms in the equations

Typical Scales

Scale	Oceanic value
L	10 km = 10^4 m
H	100 m = 10^2 m
T	≥ 1 day $\simeq 9 \times 10^4$ s
U	0.1 m/s
W	
P	variable
$\Delta\rho$	
Ω	10^{-5} s $^{-1}$.

inertial terms *rotation*

frictional forces

$$\frac{1}{\Omega T}, \quad \frac{U}{\Omega L}, \quad \frac{U}{\Omega L}, \quad \frac{WL}{UH}, \quad \frac{U}{\Omega L}$$

$$1,$$

$$\frac{P}{\rho_0 \Omega L U},$$

$$\frac{A}{\Omega L^2}, \quad \frac{\nu_E}{\Omega H^2}$$

Important Dimensionless Numbers

measuring the sizes of the terms in the equations

1 Rossby Number

advection/rotation


$$Ro = \frac{U}{\Omega L}$$

Typical Scales	
Scale	Oceanic value
L	10 km = 10^4 m
H	100 m = 10^2 m
T	≥ 1 day $\simeq 9 \times 10^4$ s
U	0.1 m/s
W	
P	variable
$\Delta\rho$	
Ω	10^{-5} s $^{-1}$.

inertial terms

rotation

frictional forces



$$\frac{1}{\Omega T}, \frac{U}{\Omega L}, \frac{U}{\Omega L}, \frac{WL}{UH}, \frac{U}{\Omega L}, \frac{1}{\Omega T}, \frac{P}{\rho_0 \Omega L U}, \frac{A}{\Omega L^2}, \frac{\nu_E}{\Omega H^2}$$

Important Dimensionless Numbers

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T	≥ 1 day $\simeq 9 \times 10^4$ s
U	0.1 m/s
W	
P	variable
$\Delta\rho$	
Ω	10^{-5} s $^{-1}$.

2 Ekman Number
viscous force/rotation

$$Ek = \frac{\nu_E}{\Omega H^2}$$

inertial terms *rotation*

frictional forces

$\frac{1}{\Omega T}$, $\frac{U}{\Omega L}$, $\frac{U}{\Omega L}$, $\frac{WL}{UH}$, $\frac{U}{\Omega L}$

1

$\frac{P}{\rho_0 \Omega L U}$, $\frac{A}{\Omega L^2}$, $\frac{\nu_E}{\Omega H^2}$

Important Dimensionless Numbers

measuring the sizes of the terms in the equations

inertial terms

frictional forces

inertial terms

rotation

frictional forces

$\frac{1}{\Omega T}$, $\frac{U}{\Omega L}$, $\frac{U}{\Omega L}$, $\frac{WL}{UH}$, $\frac{U}{\Omega L}$, 1 , $\frac{P}{\rho_0 \Omega L U}$, $\frac{A}{\Omega L^2}$, $\frac{\nu_E}{\Omega H^2}$

Important Dimensionless Numbers

measuring the sizes of the terms in the equations

3 Reynolds Number

inertial forces/frictional forces

$$Re = \frac{UL}{\nu_E} = \frac{U}{\Omega L} \cdot \frac{\Omega H^2}{\nu_E} \cdot \frac{L^2}{H^2} = \frac{Ro}{Ek} \left(\frac{L}{H} \right)^2$$

inertial terms

frictional forces

inertial terms *rotation*

frictional forces

$$\frac{1}{\Omega T}, \frac{U}{\Omega L}, \frac{U}{\Omega L}, \frac{WL}{UH} \cdot \frac{U}{\Omega L}, 1, \frac{P}{\rho_0 \Omega L U}, \frac{A}{\Omega L^2}, \frac{\nu_E}{\Omega H^2}$$

Important Dimensionless Numbers

measuring the sizes of the terms in the equations

3 Reynolds Number

inertial forces/frictional forces

high = turbulent flows

$$Re = \frac{UL}{\nu_E} = \frac{U}{\Omega L} \cdot \frac{\Omega H^2}{\nu_E} \cdot \frac{L^2}{H^2} = \frac{Ro}{Ek} \left(\frac{L}{H} \right)^2$$

inertial terms

frictional forces

inertial terms *rotation*

frictional forces

$$\frac{1}{\Omega T}, \frac{U}{\Omega L}, \frac{U}{\Omega L}, \frac{WL}{UH} \cdot \frac{U}{\Omega L}, \frac{1}{1}, \frac{P}{\rho_0 \Omega L U}, \frac{A}{\Omega L^2}, \frac{\nu_E}{\Omega H^2}$$

Important Dimensionless Numbers

measuring the sizes of the terms in the equations

inertial terms *rotation*

$$\frac{1}{\Omega T}, \quad \frac{U}{\Omega L}, \quad \frac{U}{\Omega L}, \quad \frac{WL}{UH} \cdot \frac{U}{\Omega L}$$

$$1,$$

$$\frac{P}{\rho_0 \Omega L U},$$

frictional forces

$$\frac{A}{\Omega L^2}, \quad \frac{\nu_E}{\Omega H^2}$$

Important Dimensionless Numbers

measuring the sizes of the terms in the equations

4 *Richardson Number*

inertial terms *rotation*

$$\frac{1}{\Omega T}, \quad \frac{U}{\Omega L}, \quad \frac{U}{\Omega L}, \quad \frac{WL}{UH} \cdot \frac{U}{\Omega L}$$

$$1,$$

$$\frac{P}{\rho_0 \Omega L U},$$

frictional forces

$$\frac{A}{\Omega L^2}, \quad \frac{\nu_E}{\Omega H^2}$$

Important Dimensionless Numbers

measuring the sizes of the terms in the equations

4 *Richardson Number*

$$Ri = \frac{gH\Delta\rho}{\rho_0 U^2}$$

inertial terms *rotation*

frictional forces

$\frac{1}{\Omega T}$	$\frac{U}{\Omega L}$	$\frac{U}{\Omega L}$	$\frac{WL}{UH}$	$\frac{U}{\Omega L}$	1	$\frac{P}{\rho_0 \Omega L U}$	$\frac{A}{\Omega L^2}$	$\frac{\nu_E}{\Omega H^2}$
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Important Dimensionless Numbers

measuring the sizes of the terms in the equations

4 Richardson Number

Available Potential Energy (APE)

$$Ri = \frac{gH\Delta\rho}{\rho_0 U^2}$$

Kinetic Energy (KE)

inertial terms **rotation**

frictional forces

$\frac{1}{\Omega T}$	$\frac{U}{\Omega L}$	$\frac{U}{\Omega L}$	$\frac{WL}{UH}$	$\frac{U}{\Omega L}$	1	$\frac{P}{\rho_0 \Omega L U}$	$\frac{A}{\Omega L^2}$	$\frac{\nu_E}{\Omega H^2}$
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Geostrophic flows or steady state solution

inertial terms *rotation*

$$\frac{1}{\Omega T}, \quad \frac{U}{\Omega L}, \quad \frac{U}{\Omega L}, \quad \frac{WL}{UH} \cdot \frac{U}{\Omega L}$$

$$1,$$

$$\frac{P}{\rho_0 \Omega L U},$$

frictional forces

$$\frac{A}{\Omega L^2}, \quad \frac{\nu_E}{\Omega H^2}$$

Geostrophic flows or steady state solution

$$Ro_T \ll 1, \quad Ro \ll 1, \quad Ek \ll 1.$$

inertial terms *rotation*

$$\frac{1}{\Omega T}, \quad \frac{U}{\Omega L}, \quad \frac{U}{\Omega L}, \quad \frac{WL}{UH}, \quad \frac{U}{\Omega L}$$

1,

$$\frac{P}{\rho_0 \Omega L U},$$

frictional forces

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Geostrophic flows or steady state solution

$$Ro_T \ll 1, \quad Ro \ll 1, \quad Ek \ll 1.$$

x - momentum:

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z} - fv =$$

$$-\frac{1}{\rho_0} \frac{\partial p}{\partial x} + \frac{\partial}{\partial x} \left(\mathcal{A} \frac{\partial u}{\partial x} \right) + \frac{\partial}{\partial y} \left(\mathcal{A} \frac{\partial u}{\partial y} \right) + \frac{\partial}{\partial z} \left(\nu_E \frac{\partial u}{\partial z} \right)$$

inertial terms *rotation*

frictional forces

$$\frac{1}{\Omega T}, \quad \frac{U}{\Omega L}, \quad \frac{U}{\Omega L}, \quad \frac{WL}{UH}, \quad \frac{U}{\Omega L}$$

$$1,$$

$$\frac{P}{\rho_0 \Omega L U},$$

$$\frac{\mathcal{A}}{\Omega L^2}, \quad \frac{\nu_E}{\Omega H^2}$$

Geostrophic flows or steady state solution

$$Ro_T \ll 1, \quad Ro \ll 1, \quad Ek \ll 1.$$

x - momentum:

$$\begin{aligned} & \cancel{\frac{\partial u}{\partial t}} + u \cancel{\frac{\partial u}{\partial x}} + v \cancel{\frac{\partial u}{\partial y}} + w \cancel{\frac{\partial u}{\partial z}} - fv = \\ & -\frac{1}{\rho_0} \frac{\partial p}{\partial x} + \cancel{\frac{\partial}{\partial x} \left(\mathcal{A} \frac{\partial u}{\partial x} \right)} + \cancel{\frac{\partial}{\partial y} \left(\mathcal{A} \frac{\partial u}{\partial y} \right)} + \cancel{\frac{\partial}{\partial z} \left(\nu_E \frac{\partial u}{\partial z} \right)} \end{aligned}$$

inertial terms

rotation

frictional forces

$$\frac{1}{\Omega T}, \quad \frac{U}{\Omega L}, \quad \frac{U}{\Omega L}, \quad \frac{WL}{UH}, \quad \frac{U}{\Omega L}$$

$$1,$$

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Geostrophic flows or steady state solution

$$Ro_T \ll 1, \quad Ro \ll 1, \quad Ek \ll 1.$$

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$$\frac{1}{\rho_0} \frac{\partial p}{\partial x} + \frac{\partial}{\partial x} \left(\cancel{A \frac{\partial u}{\partial x}} \right) + \frac{\partial}{\partial y} \left(\cancel{A \frac{\partial u}{\partial y}} \right) + \frac{\partial}{\partial z} \left(\cancel{\nu_E \frac{\partial u}{\partial z}} \right)$$

inertial terms

rotation

frictional forces

$$\frac{1}{\Omega T}, \quad \frac{U}{\Omega L}, \quad \frac{U}{\Omega L}, \quad \frac{WL}{UH}, \quad \frac{U}{\Omega L}$$

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Geostrophic flows or steady state solution

$$Ro_T \ll 1, \quad Ro \ll 1, \quad Ek \ll 1.$$

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$$\cancel{\frac{\partial u}{\partial t}} + u \cancel{\frac{\partial u}{\partial x}} + v \cancel{\frac{\partial u}{\partial y}} + w \cancel{\frac{\partial u}{\partial z}} - fv =$$

$$-\frac{1}{\rho_0} \frac{\partial p}{\partial x} + \frac{\partial}{\partial x} \left(\mathcal{A} \frac{\partial u}{\partial x} \right) + \frac{\partial}{\partial y} \left(\mathcal{A} \frac{\partial u}{\partial y} \right) + \frac{\partial}{\partial z} \left(\nu_E \frac{\partial u}{\partial z} \right)$$

small small small small small small small

z – momentum:

$$0 = -\frac{\partial p}{\partial z} - \rho g$$

continuity:

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0$$

energy:

$$\frac{\partial \rho}{\partial t} + u \frac{\partial \rho}{\partial x} + v \frac{\partial \rho}{\partial y} + w \frac{\partial \rho}{\partial z} =$$

$$\frac{\partial}{\partial x} \left(\mathcal{A} \frac{\partial \rho}{\partial x} \right) + \frac{\partial}{\partial y} \left(\mathcal{A} \frac{\partial \rho}{\partial y} \right) + \frac{\partial}{\partial z} \left(\kappa_E \frac{\partial \rho}{\partial z} \right)$$

Geostrophic flows or steady state solution

$$Ro_T \ll 1, \quad Ro \ll 1, \quad Ek \ll 1.$$

$$\rho = 0 \text{ (no density variation)}$$

x – momentum:

$$\cancel{\frac{\partial u}{\partial t}} + u \cancel{\frac{\partial u}{\partial x}} + v \cancel{\frac{\partial u}{\partial y}} + w \cancel{\frac{\partial u}{\partial z}} - \boxed{fv} =$$

$$\boxed{-\frac{1}{\rho_0} \frac{\partial p}{\partial x}} + \frac{\partial}{\partial x} \left(\cancel{\mathcal{A} \frac{\partial u}{\partial x}} \right) + \frac{\partial}{\partial y} \left(\cancel{\mathcal{A} \frac{\partial u}{\partial y}} \right) + \frac{\partial}{\partial z} \left(\cancel{\nu_E \frac{\partial u}{\partial z}} \right)$$

small small small small small small small

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$$\frac{\partial \rho}{\partial t} + u \frac{\partial \rho}{\partial x} + v \frac{\partial \rho}{\partial y} + w \frac{\partial \rho}{\partial z} =$$

$$\frac{\partial}{\partial x} \left(\mathcal{A} \frac{\partial \rho}{\partial x} \right) + \frac{\partial}{\partial y} \left(\mathcal{A} \frac{\partial \rho}{\partial y} \right) + \frac{\partial}{\partial z} \left(\kappa_E \frac{\partial \rho}{\partial z} \right)$$

Geostrophic flows or steady state solution

$$Ro_T \ll 1, \quad Ro \ll 1, \quad Ek \ll 1.$$

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$$\boxed{-\frac{1}{\rho_0} \frac{\partial p}{\partial x}} + \frac{\partial}{\partial x} \left(\cancel{A \frac{\partial u}{\partial x}} \right) + \frac{\partial}{\partial y} \left(\cancel{A \frac{\partial u}{\partial y}} \right) + \frac{\partial}{\partial z} \left(\cancel{\nu_E \frac{\partial u}{\partial z}} \right)$$

small small small small small small small

z – momentum:

$$0 = \cancel{\frac{\partial p}{\partial z}} - \cancel{\rho g} \text{ vanish}$$

continuity:

$$\boxed{\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0}$$

energy:

$$\cancel{\frac{\partial \rho}{\partial t} + u \frac{\partial \rho}{\partial x} + v \frac{\partial \rho}{\partial y} + w \frac{\partial \rho}{\partial z} =}$$

$$\cancel{\frac{\partial}{\partial x} \left(A \frac{\partial \rho}{\partial x} \right) + \frac{\partial}{\partial y} \left(A \frac{\partial \rho}{\partial y} \right) + \frac{\partial}{\partial z} \left(\cancel{\kappa_E \frac{\partial \rho}{\partial z}} \right)} \text{ vanish}$$

Geostrophic flows or steady state solution

$$Ro_T \ll 1, \quad Ro \ll 1, \quad Ek \ll 1.$$
$$\rho = 0 \text{ (no density variation)}$$

$$-fv = -\frac{1}{\rho_0} \frac{\partial p}{\partial x}$$

$$+fu = -\frac{1}{\rho_0} \frac{\partial p}{\partial y}$$

$$0 = -\frac{1}{\rho_0} \frac{\partial p}{\partial z}$$

homogeneous flow

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0,$$

Properties of Geostrophic **homogeneous** flows

$$u = \frac{-1}{\rho_0 f} \frac{\partial p}{\partial y}, \quad v = \frac{+1}{\rho_0 f} \frac{\partial p}{\partial x}$$

Properties of Geostrophic **homogeneous** flows

geostrophic flows follow lines of constant pressure = **isobars**

$$u = \frac{-1}{\rho_0 f} \frac{\partial p}{\partial y}, \quad v = \frac{+1}{\rho_0 f} \frac{\partial p}{\partial x}$$

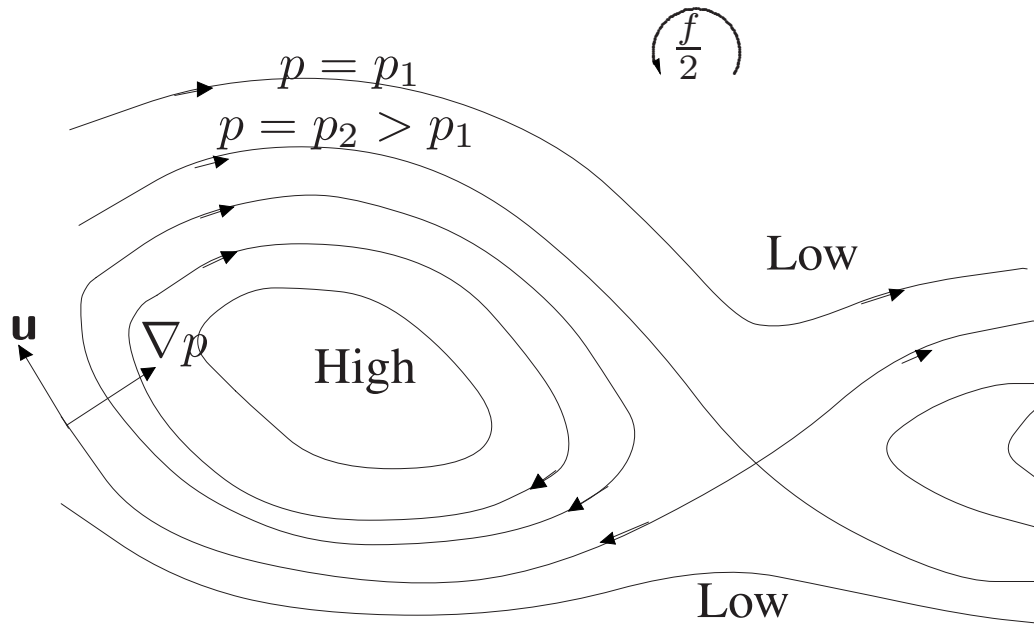


Figure 7-1 Example of geostrophic flow. The velocity vector is everywhere parallel to the lines of equal pressure. Thus, pressure contours act as streamlines. In the Northern Hemisphere (as pictured here), the fluid circulates with the high pressure on its right. The opposite holds for the Southern Hemisphere.

Properties of Geostrophic **homogeneous** flows

$$-fv = -\frac{1}{\rho_0} \frac{\partial p}{\partial x}$$

$$+fu = -\frac{1}{\rho_0} \frac{\partial p}{\partial y}$$

$$0 = -\frac{1}{\rho_0} \frac{\partial p}{\partial z}$$

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0,$$

Properties of Geostrophic **homogeneous** flows

$$\frac{\partial}{\partial z} \left(\begin{array}{rcl} -fv & = & -\frac{1}{\rho_0} \frac{\partial p}{\partial x} \\ +fu & = & -\frac{1}{\rho_0} \frac{\partial p}{\partial y} \\ 0 & = & -\frac{1}{\rho_0} \frac{\partial p}{\partial z} \end{array} \right)$$
$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0,$$

Properties of Geostrophic **homogeneous** flows

$$\frac{\partial}{\partial z} \left(\begin{array}{rcl} -fv & = & -\frac{1}{\rho_0} \frac{\partial p}{\partial x} \\ +fu & = & -\frac{1}{\rho_0} \frac{\partial p}{\partial y} \\ 0 & = & -\frac{1}{\rho_0} \frac{\partial p}{\partial z} \end{array} \right)$$
$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0,$$

no vertical shear

$$-f \frac{\partial v}{\partial z} = -\frac{1}{\rho_0} \frac{\partial}{\partial z} \left(\frac{\partial p}{\partial x} \right) = -\frac{1}{\rho_0} \frac{\partial}{\partial x} \left(\frac{\partial p}{\partial z} \right) = 0$$

Properties of Geostrophic **homogeneous** flows

$$\frac{\partial}{\partial z} \left(\begin{array}{rcl} -fv & = & -\frac{1}{\rho_0} \frac{\partial p}{\partial x} \\ +fu & = & -\frac{1}{\rho_0} \frac{\partial p}{\partial y} \\ 0 & = & -\frac{1}{\rho_0} \frac{\partial p}{\partial z} \end{array} \right)$$
$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0,$$

no vertical shear

$$-f \frac{\partial v}{\partial z} = -\frac{1}{\rho_0} \frac{\partial}{\partial z} \left(\frac{\partial p}{\partial x} \right) = -\frac{1}{\rho_0} \frac{\partial}{\partial x} \left(\frac{\partial p}{\partial z} \right) = 0$$

Taylor–Proudman theorem (Taylor, 1923; Proudman, 1953)

Properties of Geostrophic **homogeneous** flows

$$-fv = -\frac{1}{\rho_0} \frac{\partial p}{\partial x}$$

$$+fu = -\frac{1}{\rho_0} \frac{\partial p}{\partial y}$$

$$0 = -\frac{1}{\rho_0} \frac{\partial p}{\partial z}$$

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0,$$

Properties of Geostrophic **homogeneous** flows

if f-plane

$$-fv = -\frac{1}{\rho_0} \frac{\partial p}{\partial x}$$

$$+fu = -\frac{1}{\rho_0} \frac{\partial p}{\partial y}$$

$$0 = -\frac{1}{\rho_0} \frac{\partial p}{\partial z}$$

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0,$$

Properties of Geostrophic **homogeneous** flows

if f-plane

$$\frac{\partial}{\partial x} \left(\begin{array}{l} -fv = -\frac{1}{\rho_0} \frac{\partial p}{\partial x} \\ +fu = -\frac{1}{\rho_0} \frac{\partial p}{\partial y} \\ 0 = -\frac{1}{\rho_0} \frac{\partial p}{\partial z} \end{array} \right)$$
$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0,$$

Properties of Geostrophic **homogeneous** flows

if f-plane

$$\frac{\partial}{\partial x} \left(\begin{array}{l} -fv = -\frac{1}{\rho_0} \frac{\partial p}{\partial x} \\ +fu = -\frac{1}{\rho_0} \frac{\partial p}{\partial y} \\ 0 = -\frac{1}{\rho_0} \frac{\partial p}{\partial z} \end{array} \right)$$
$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0,$$

non-divergent



$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} = -\frac{\partial}{\partial x} \left(\frac{1}{\rho_0 f} \frac{\partial p}{\partial y} \right) + \frac{\partial}{\partial y} \left(\frac{1}{\rho_0 f} \frac{\partial p}{\partial x} \right) = 0$$

Properties of Geostrophic **homogeneous** flows

if f-plane

$$\frac{\partial}{\partial x} \left(\begin{array}{rcl} -fv & = & -\frac{1}{\rho_0} \frac{\partial p}{\partial x} \\ +fu & = & -\frac{1}{\rho_0} \frac{\partial p}{\partial y} \\ 0 & = & -\frac{1}{\rho_0} \frac{\partial p}{\partial z} \end{array} \right)$$

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0,$$

non-divergent



$$\frac{\partial w}{\partial z} = 0$$

Properties of Geostrophic **homogeneous** flows

if f -plane

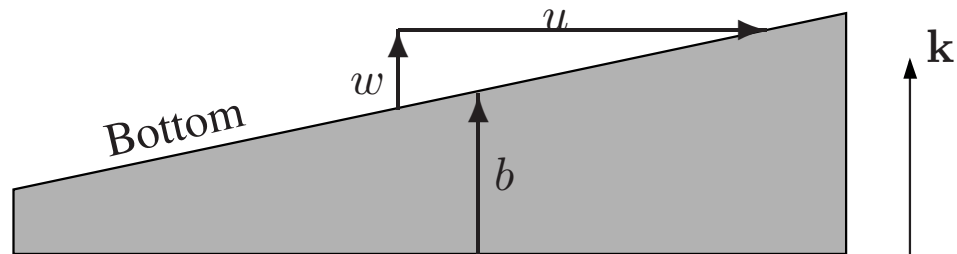


Figure 7-3 Schematic view of a flow over a sloping bottom. A vertical velocity must accompany flow across isobaths.

$$\frac{\partial w}{\partial z} = 0$$

Properties of Geostrophic **homogeneous** flows

if f -plane

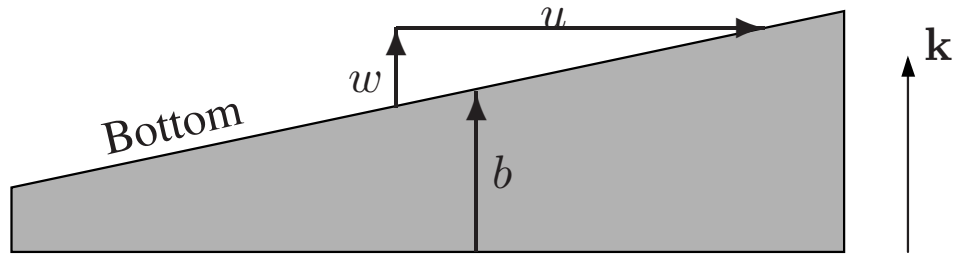


Figure 7-3 Schematic view of a flow over a sloping bottom. A vertical velocity must accompany flow across isobaths.

therefore flow is permitted only along isobaths

$$\frac{\partial w}{\partial z} = 0$$

Properties of Geostrophic **homogeneous** flows

if f-plane

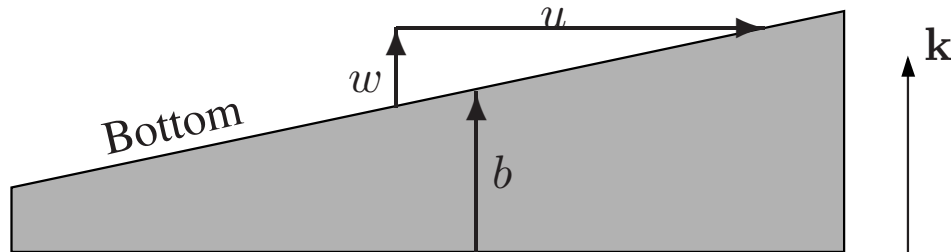
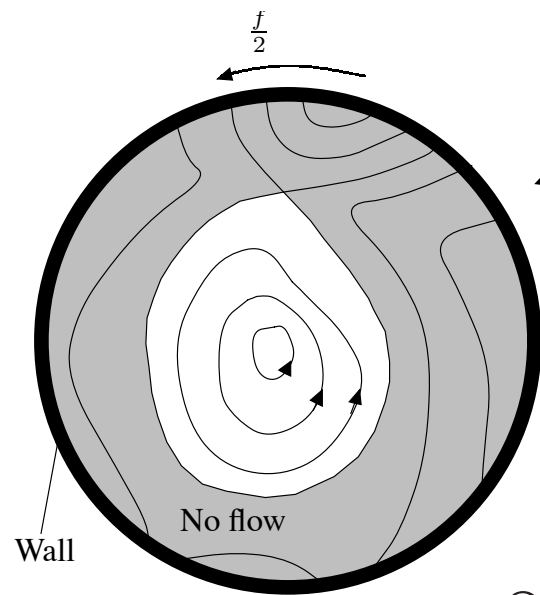


Figure 7-3 Schematic view of a flow over a sloping bottom. A vertical velocity must accompany flow across isobaths.



therefore flow is permitted only along isobaths

Figure 7-4 Geostrophic flow in a closed domain and over irregular topography. Solid lines are isobaths (contours of equal depth). Flow is permitted only along closed isobaths

$$\frac{\partial w}{\partial z} = 0$$

Properties of Geostrophic **homogeneous** flows

$$-fv = -\frac{1}{\rho_0} \frac{\partial p}{\partial x}$$

$$+fu = -\frac{1}{\rho_0} \frac{\partial p}{\partial y}$$

$$0 = -\frac{1}{\rho_0} \frac{\partial p}{\partial z}$$

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0,$$

homogeneous flow

Properties of Geostrophic **homogeneous** flows

$$-fv = -\frac{1}{\rho_0} \frac{\partial p}{\partial x}$$

$$+fu = -\frac{1}{\rho_0} \frac{\partial p}{\partial y}$$

$$0 = -\frac{1}{\rho_0} \frac{\partial p}{\partial z}$$

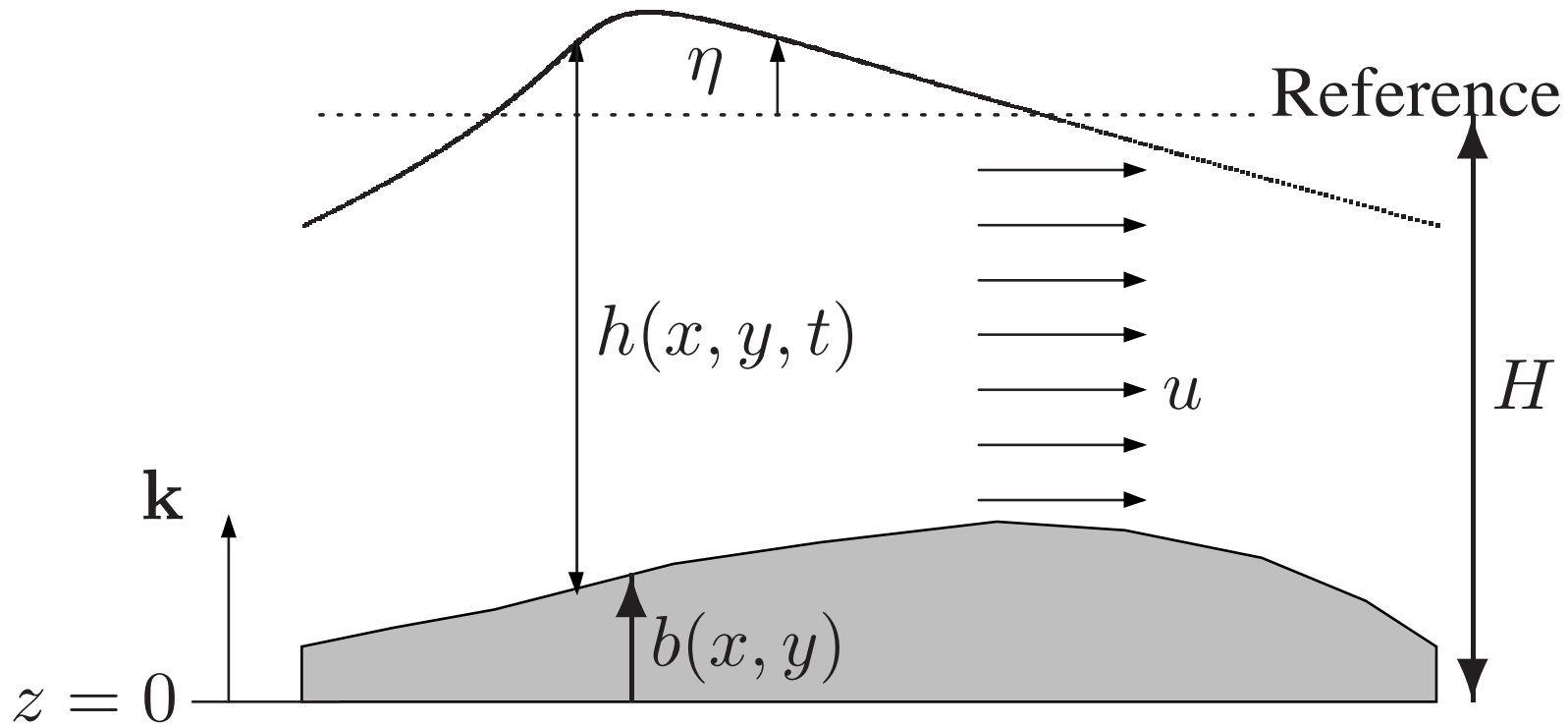
homogeneous flow

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0,$$

if pressure only a function of the sea level

$$p = \rho_0 g \eta$$

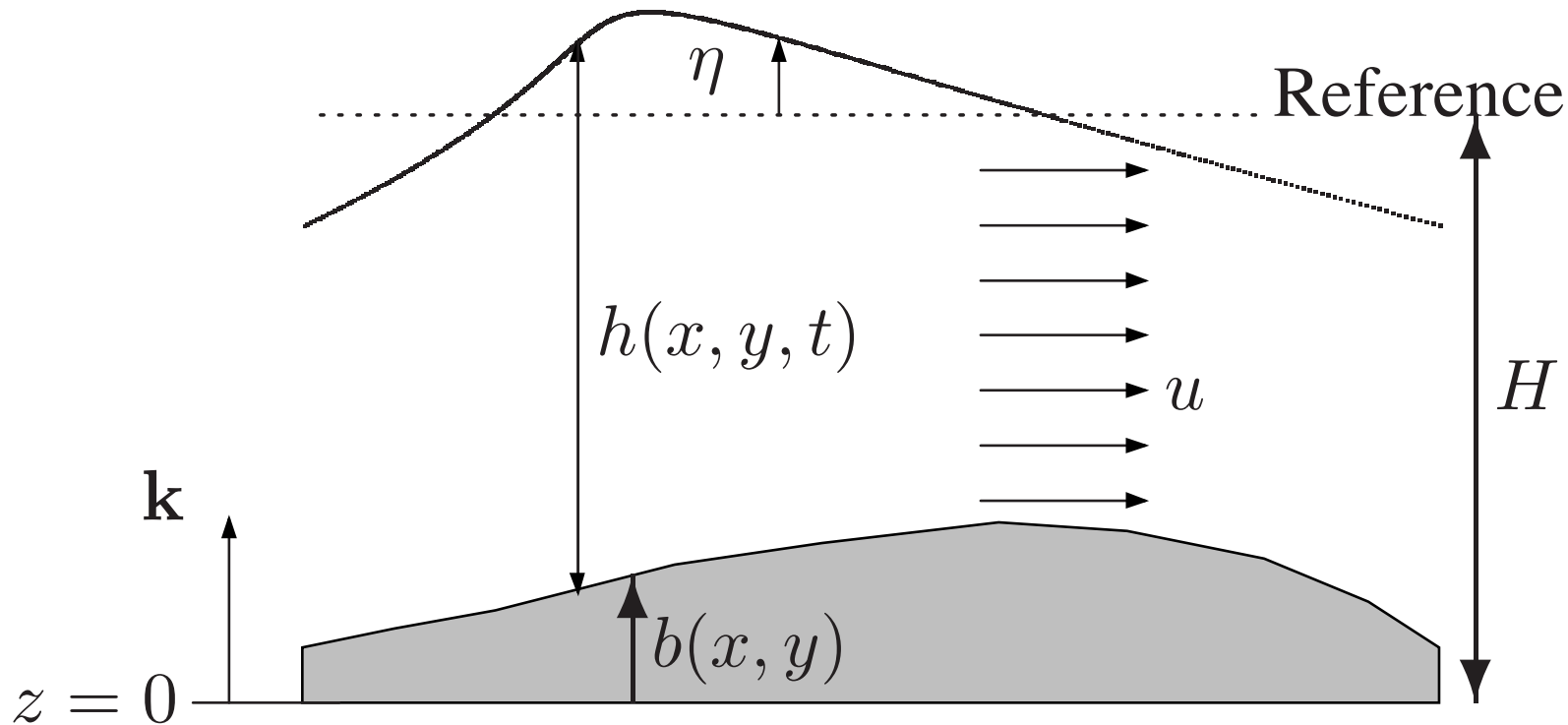
Properties of Geostrophic **homogeneous** flows



if pressure only a function of the sea level

$$p = \rho_0 g \eta$$

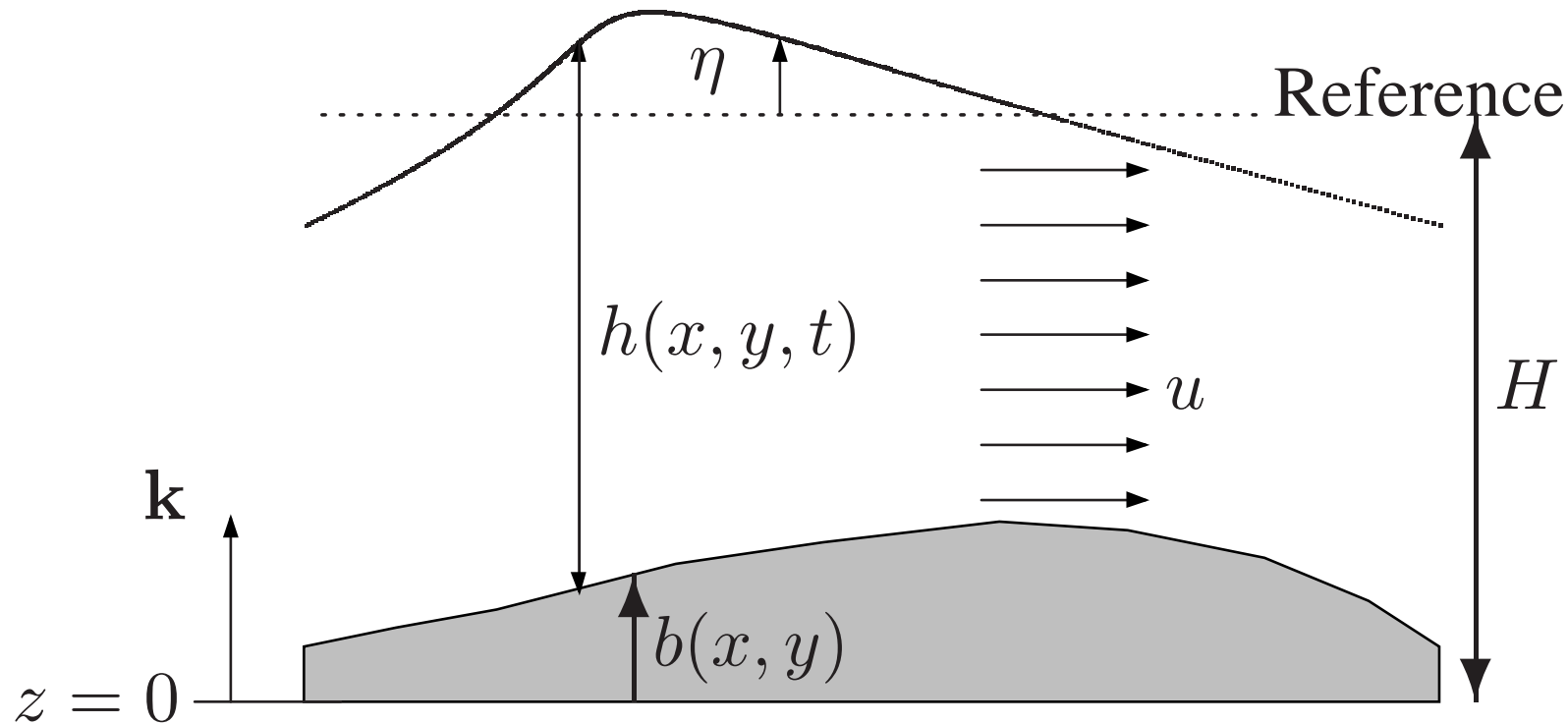
Properties of Geostrophic **homogeneous** flows



re-write the continuity equation for the layer velocity

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0$$

Properties of Geostrophic **homogeneous** flows



re-write the continuity equation for the layer velocity

$$\left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0 \right) \int_b^{b+h} dz$$

integrate from bottom to surface

Properties of Geostrophic **homogeneous** flows

re-write the continuity equation for the layer velocity

$$\left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} \right) \int_b^{b+h} dz + [w]_b^{b+h}$$

integrate from bottom to surface

Properties of Geostrophic **homogeneous** flows

$$\begin{aligned}w(z = b + h) &= \frac{\partial}{\partial t}(b + h) + u \frac{\partial}{\partial x}(b + h) + v \frac{\partial}{\partial y}(b + h) \\ &= \frac{\partial \eta}{\partial t} + u \frac{\partial \eta}{\partial x} + v \frac{\partial \eta}{\partial y} \\ w(z = b) &= u \frac{\partial b}{\partial x} + v \frac{\partial b}{\partial y} .\end{aligned}$$

boundary conditions

re-write the continuity equation for the layer velocity

$$\left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} \right) \int_b^{b+h} dz + [w]_b^{b+h}$$

integrate from bottom to surface

Properties of Geostrophic **homogeneous** flows

$$\begin{aligned}w(z = b + h) &= \frac{\partial}{\partial t}(b + h) + u \frac{\partial}{\partial x}(b + h) + v \frac{\partial}{\partial y}(b + h) \\ &= \frac{\partial \eta}{\partial t} + u \frac{\partial \eta}{\partial x} + v \frac{\partial \eta}{\partial y} \\ w(z = b) &= u \frac{\partial b}{\partial x} + v \frac{\partial b}{\partial y} .\end{aligned}$$

boundary conditions

where $\eta = b + h - H$

re-write the continuity equation for the layer velocity

$$\left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} \right) \int_b^{b+h} dz + [w]_b^{b+h}$$

integrate from bottom to surface

Properties of Geostrophic **homogeneous** flows

$$\begin{aligned}w(z = b + h) &= \frac{\partial}{\partial t}(b + h) + u \frac{\partial}{\partial x}(b + h) + v \frac{\partial}{\partial y}(b + h) \\ &= \frac{\partial \eta}{\partial t} + u \frac{\partial \eta}{\partial x} + v \frac{\partial \eta}{\partial y} \\ w(z = b) &= u \frac{\partial b}{\partial x} + v \frac{\partial b}{\partial y} .\end{aligned}$$

boundary conditions

where $\eta = b + h - H$

re-write the continuity equation for the layer velocity

$$\frac{\partial \eta}{\partial t} + \frac{\partial}{\partial x}(hu) + \frac{\partial}{\partial y}(hv) = 0;$$

NEW continuity equation!

Equations of Geostrophic **homogeneous** flows

$$-fv = -\frac{1}{\rho_0} \frac{\partial p}{\partial x}$$

$$+fu = -\frac{1}{\rho_0} \frac{\partial p}{\partial y}$$

$$\frac{\partial \eta}{\partial t} + \frac{\partial}{\partial x}(hu) + \frac{\partial}{\partial y}(hv) = 0;$$

Equations of Geostrophic **homogeneous** flows

if pressure only a function of the sea level

$$p = \rho_0 g \eta$$

$$\begin{aligned} -fv &= -\frac{1}{\rho_0} \frac{\partial p}{\partial x} \\ +fu &= -\frac{1}{\rho_0} \frac{\partial p}{\partial y} \end{aligned}$$

$$\frac{\partial \eta}{\partial t} + \frac{\partial}{\partial x}(hu) + \frac{\partial}{\partial y}(hv) = 0;$$

Equations of Geostrophic **homogeneous** flows

$$-fv = -g \frac{\partial \eta}{\partial x}$$

$$+fu = -g \frac{\partial \eta}{\partial y}$$

$$\frac{\partial \eta}{\partial t} + \frac{\partial}{\partial x}(hu) + \frac{\partial}{\partial y}(hv) = 0;$$

Equations of Geostrophic **homogeneous** flows

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} - fv = -g \frac{\partial \eta}{\partial x}$$

$$\frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} + fu = -g \frac{\partial \eta}{\partial y}$$

$$\frac{\partial \eta}{\partial t} + \frac{\partial}{\partial x}(hu) + \frac{\partial}{\partial y}(hv) = 0,$$

Equations of Geostrophic **homogeneous** flows

shallow-water model or *barotropic equations*

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} - fv = -g \frac{\partial \eta}{\partial x}$$

$$\frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} + fu = -g \frac{\partial \eta}{\partial y}$$

$$\frac{\partial \eta}{\partial t} + \frac{\partial}{\partial x}(hu) + \frac{\partial}{\partial y}(hv) = 0;$$

Equations of Geostrophic **homogeneous** flows

shallow-water model or *barotropic equations*

describe unsteady motions of a 2D uniform density layer

or

of the depth average motion

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} - fv = -g \frac{\partial \eta}{\partial x}$$

$$\frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} + fu = -g \frac{\partial \eta}{\partial y}$$

$$\frac{\partial \eta}{\partial t} + \frac{\partial}{\partial x}(hu) + \frac{\partial}{\partial y}(hv) = 0,$$