

Shallow-water environments and their sediments

Shelf Seas: areas of submerged continental crust (submerged continental shelf)

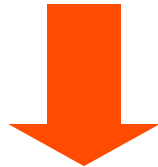
TYPES OF SEDIMENTS

Sediments in shallow-water are **TERRIGENOUS**
(**BIOGENIC**)

Weathering

Physical (or mechanical)

Chemical



Rock fragments

Quartz

Clays (2 μm)

PROCESSES AFFECTING SEDIMENTATION

Oceanic Transport

Weathering

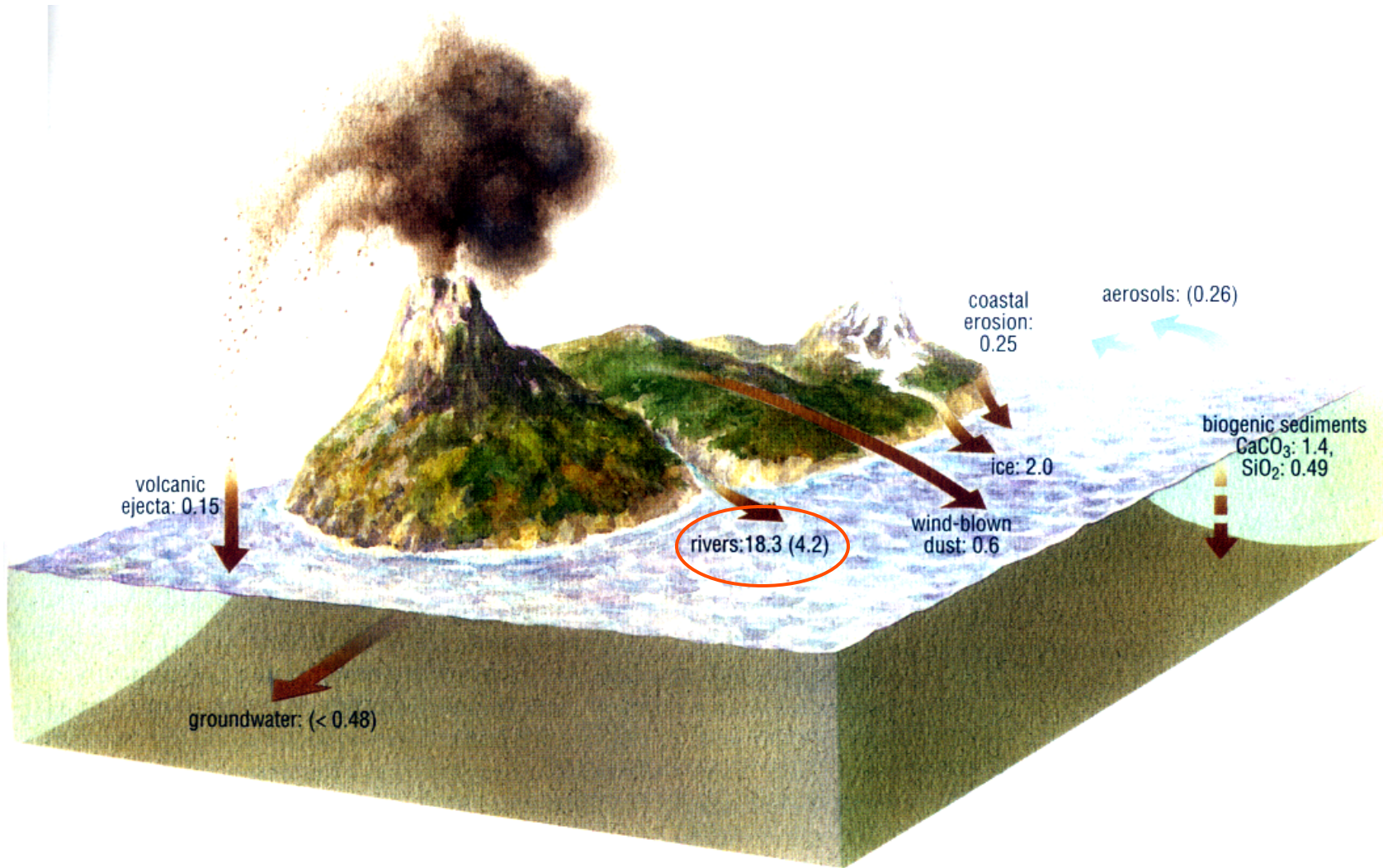
**Physical (or mechanical)
Chemical**

**Climate
(latitude)**

Topography

SOURCES/SUPPLY OF SEDIMENTS

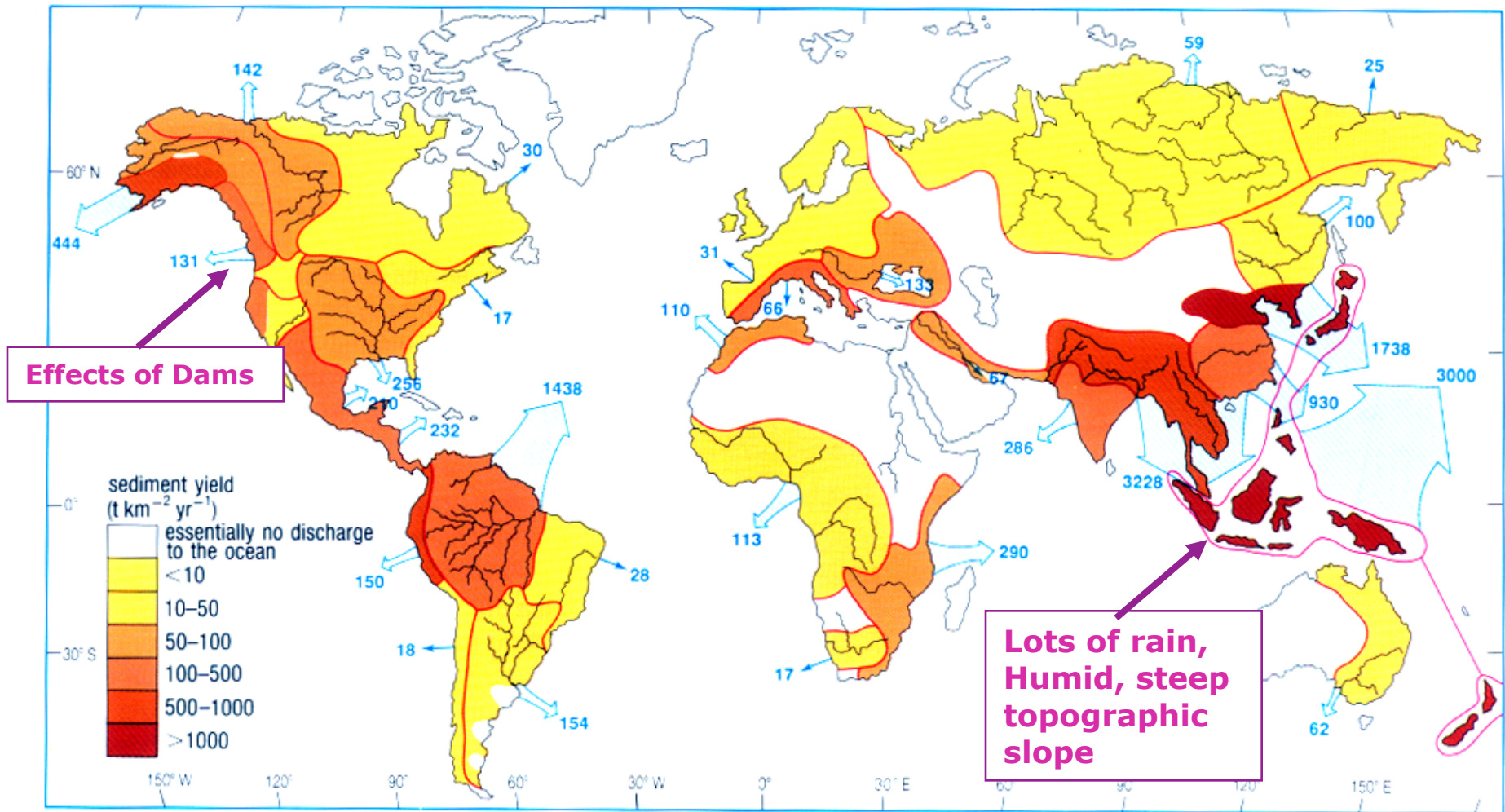
Supply of Sediments to Shelf Seas



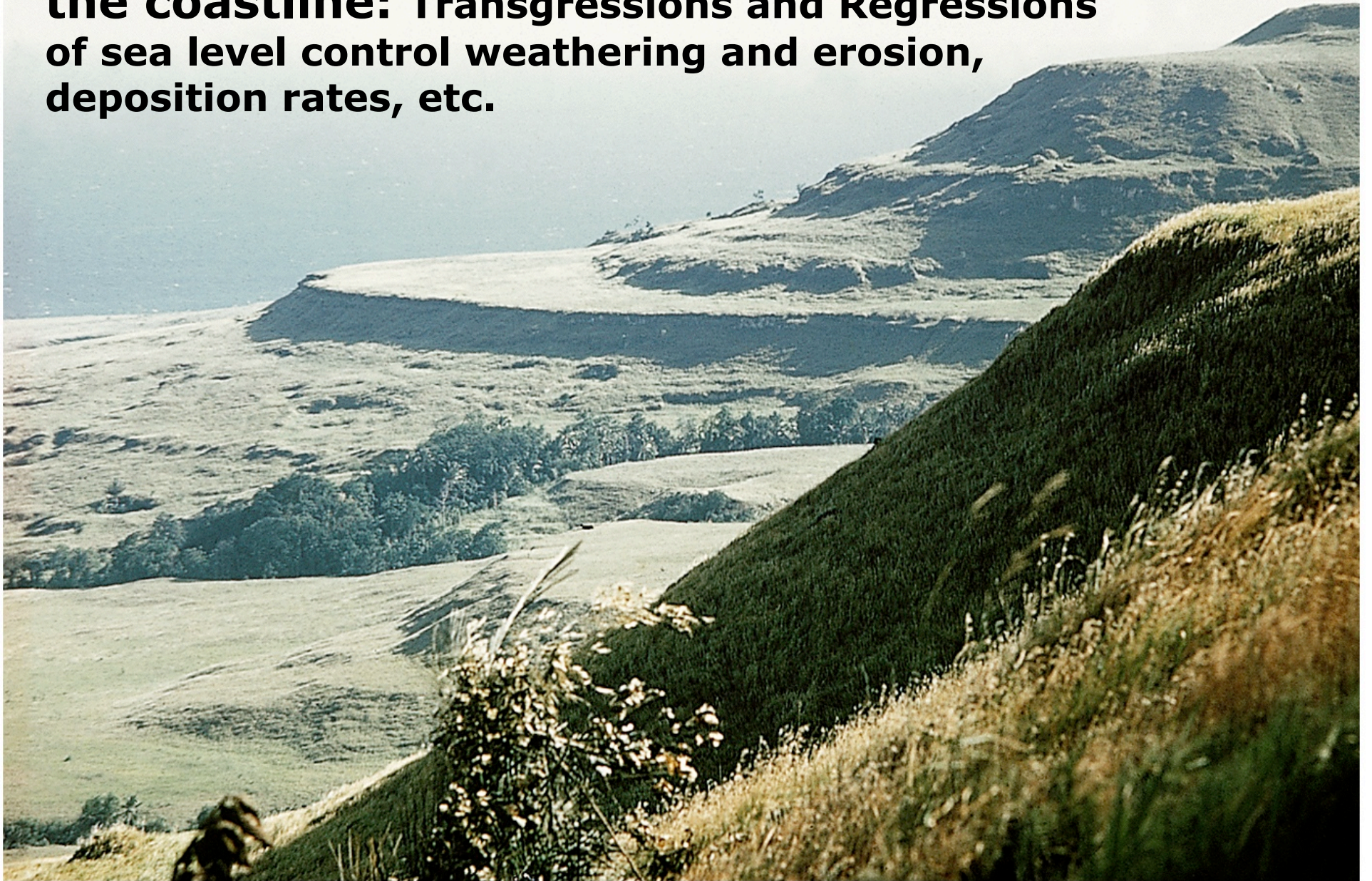
UNITS: 10⁹ tonnes/year

SPATIAL DISTRIBUTIONS of SEDIMENT SUPPLY

Discharge from major drainage basins



On long time scales ICE AGES modulate the coastline: Transgressions and Regressions of sea level control weathering and erosion, deposition rates, etc.



Shelf sediments - summary

TYPES

Rock fragments, Quartz and Clays

SOURCE (MAJOR)

Globally: Rivers

Locally: Ice and wind transport and volcanic eruption

SHAPE/STRUCTURE of SHELF SEAS

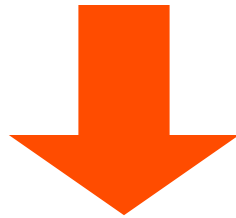
Sea level changes: deposition of rivers/glacial sediments → relict sediments



Reworked by waves and tidal currents

Our GOAL is to understand:

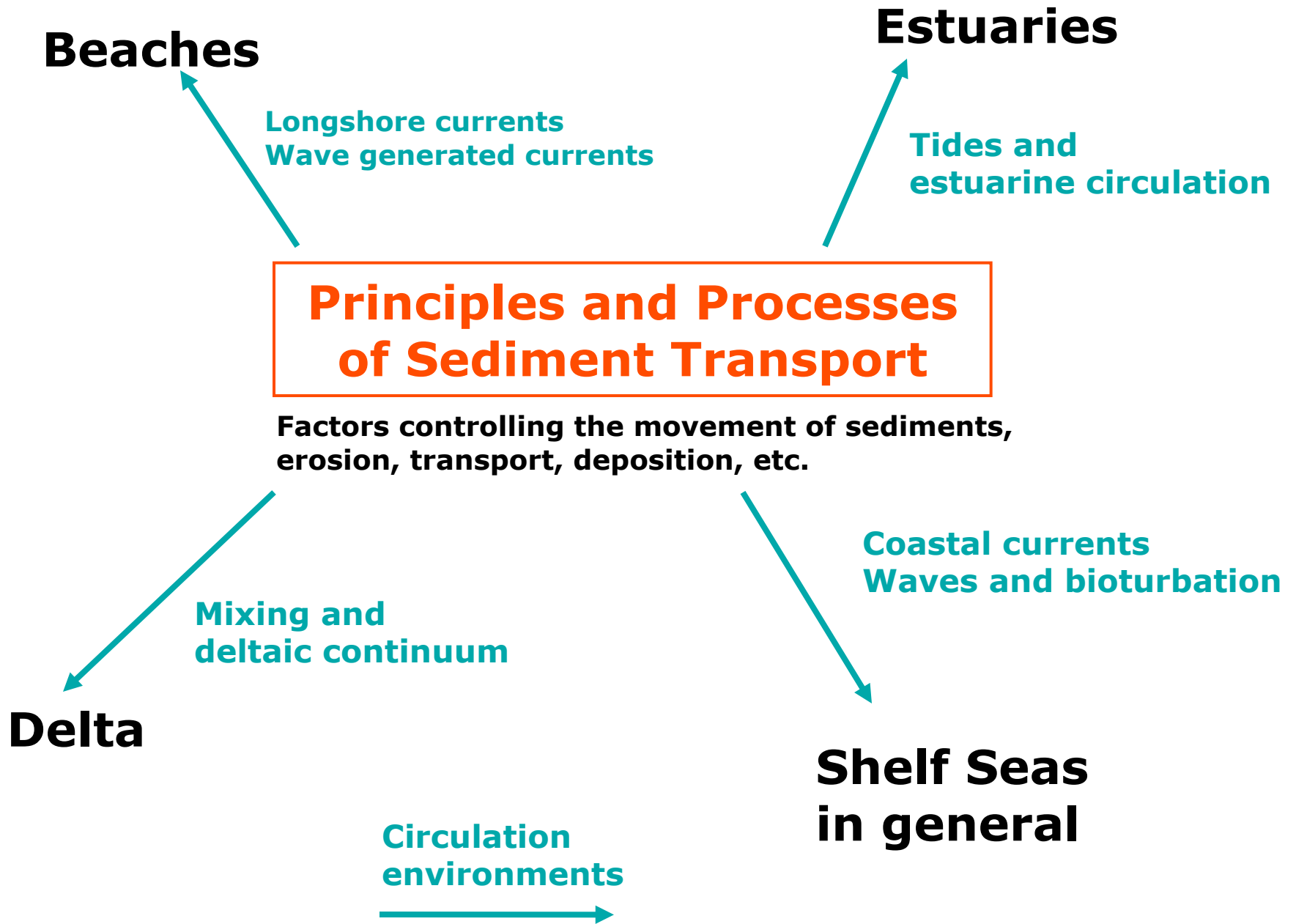
Physical and chemical principles that underlie the transport and deposition of sediments in the near shore.



Protect and Manage coastal environments from human intervention

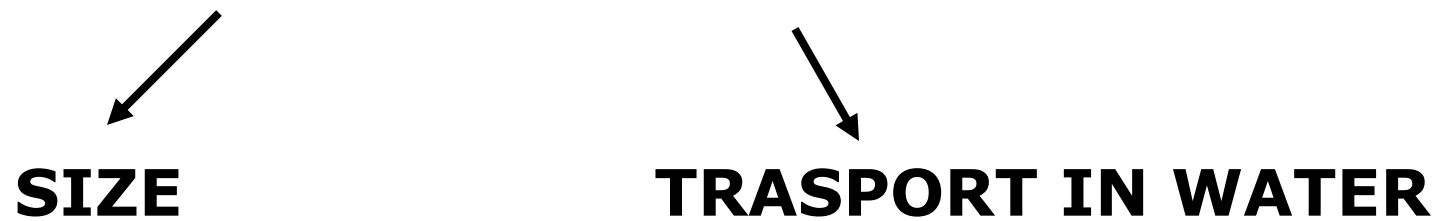
Important principle:

Equilibrium between sedimentation rate and redistribution of sediments

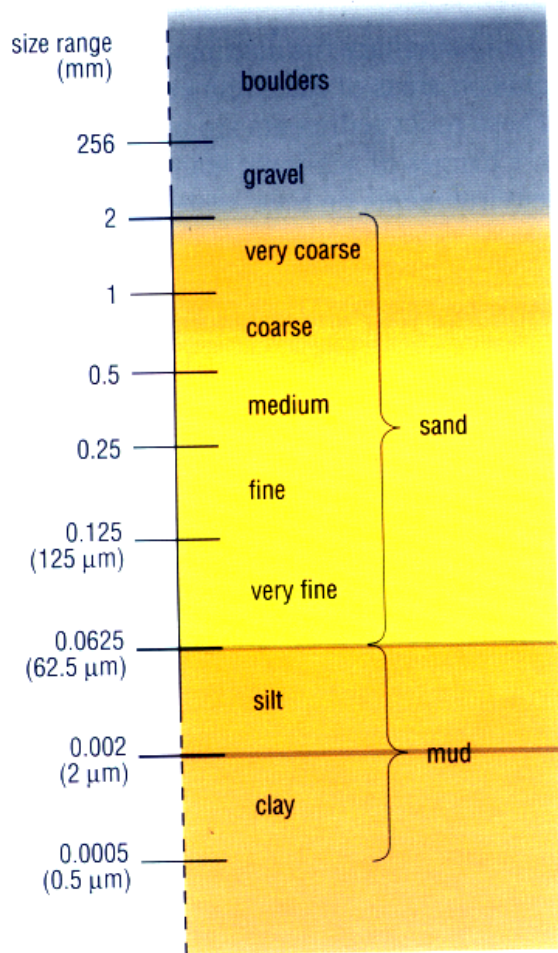


Principles and Processes of Sediment Transport

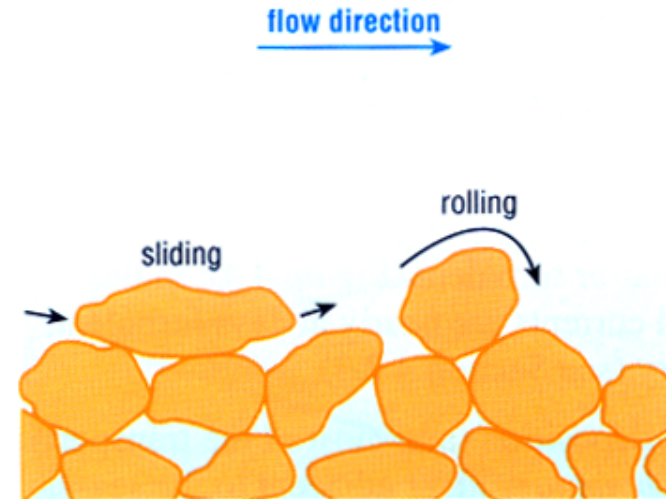
Factors controlling the movement of sediments



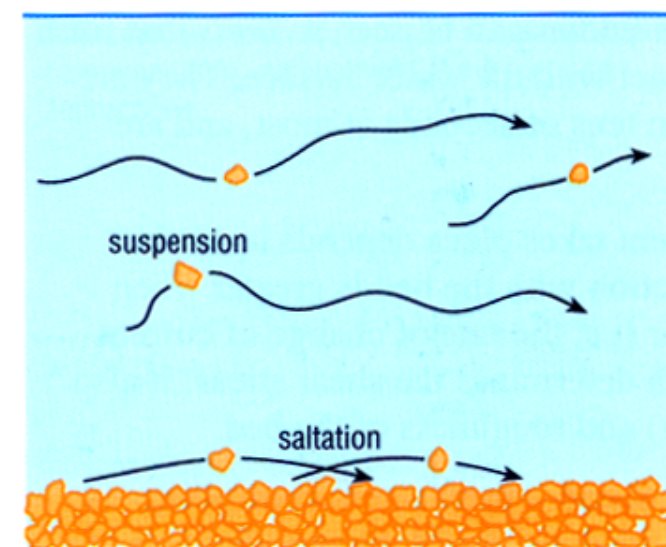
SIZE Classification



TRANSPORT in WATER



(a)



(b)

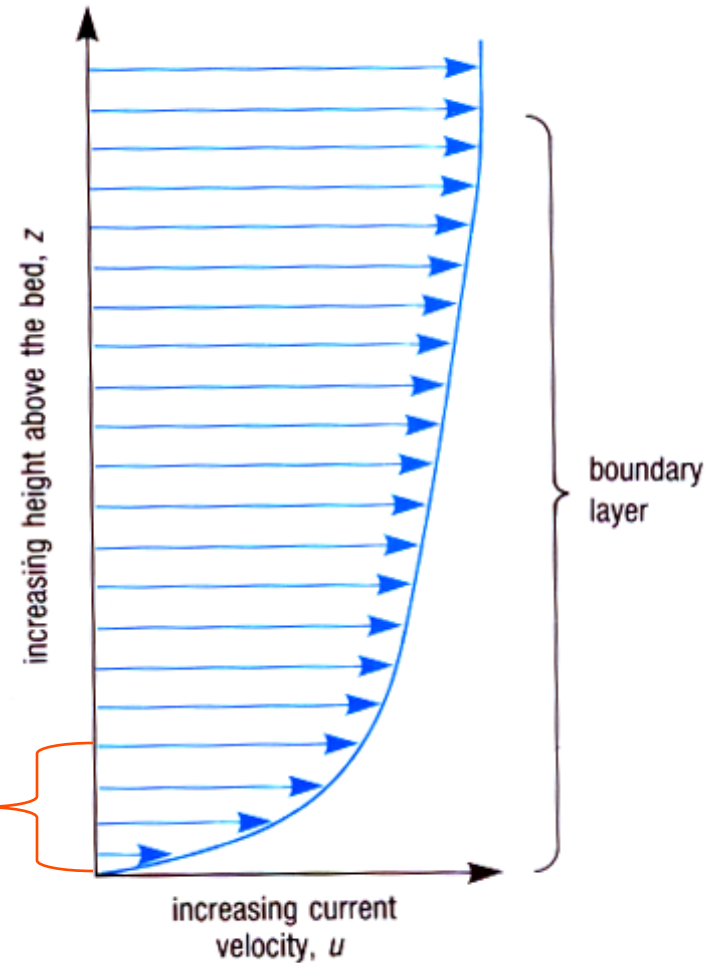
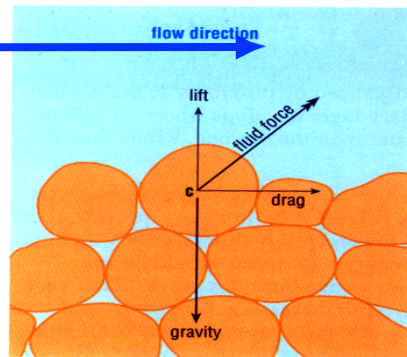
Boundary Layer and Current Shear

Movement of sediments in the **Benthic Boundary Layer** depends on:

Roughness of sediment bed
Density contrast
Degree of Turbulence
Current Shear
Shear Stress

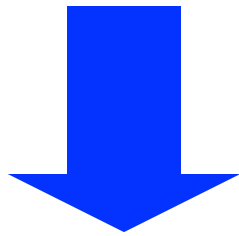
Benthic Boundary Layer
several 10s of meters

**Laminar
or
Turbulent**

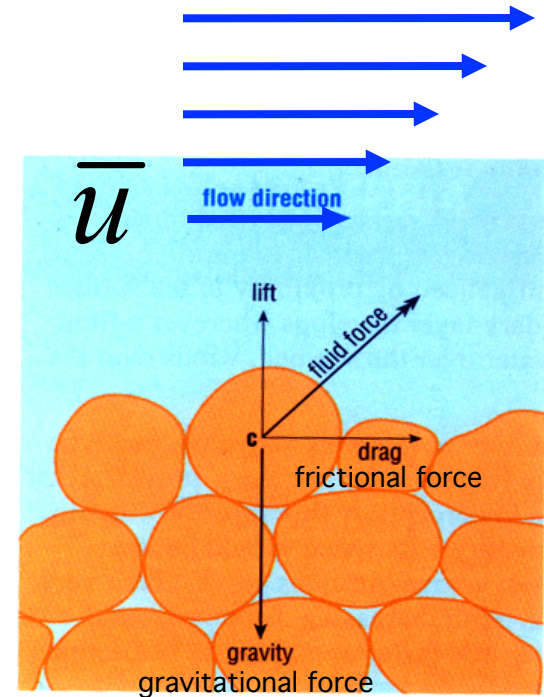


Shear Stress is proportional to velocity at the bed

$$\tau = (\mu + \eta) \frac{d\bar{u}}{dz} = \rho u_*^2$$



when strong to overcome
frictional and **gravitational** forces
Critical Shear Stress

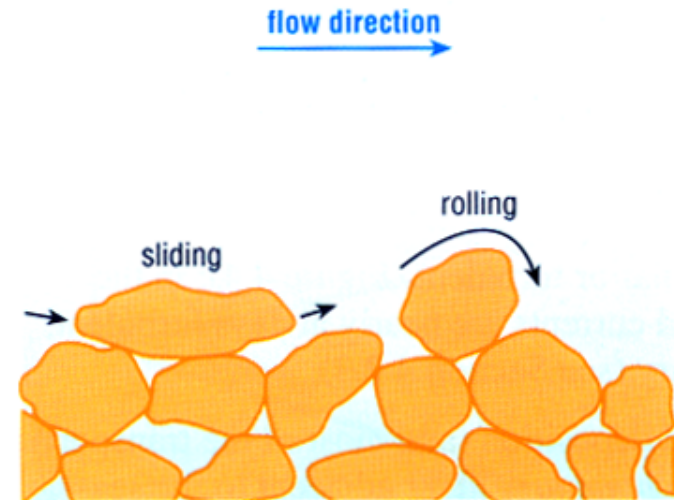


TRANSPORT in WATER

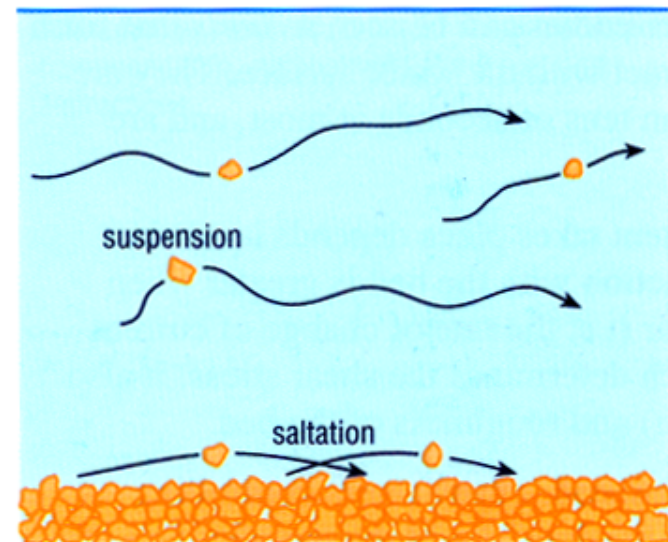
Erosion: the process by which we set sediments in motion.

Also take into account:

COHESIVE vs. NON-COHESIVE



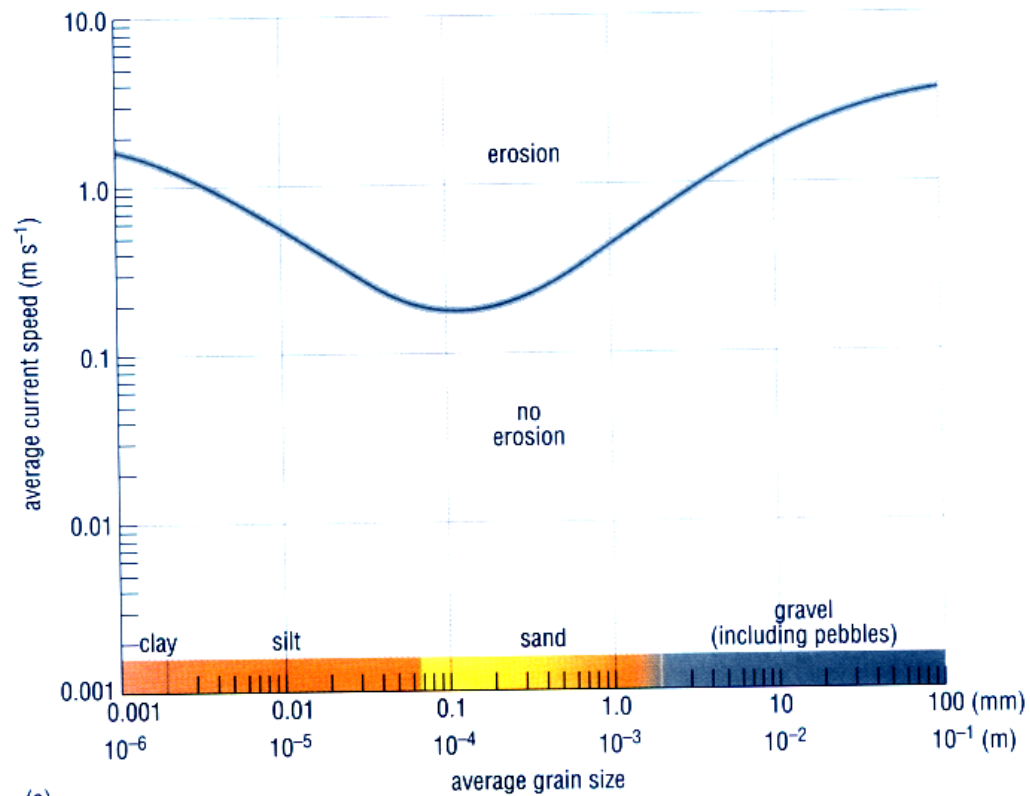
(a)



(b)

WHEN DOES EROSION OCCUR?

\bar{u}

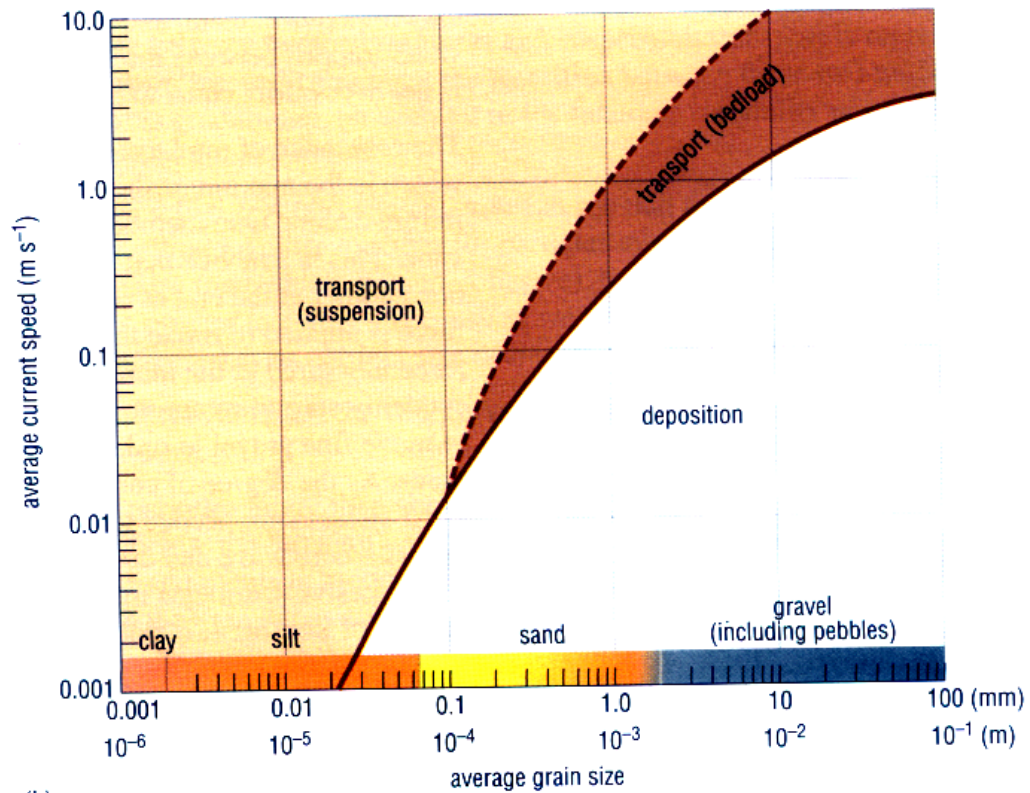


(a)

Size of grain

TYPES of TRANSPORT EXPECTED and DEPOSITION

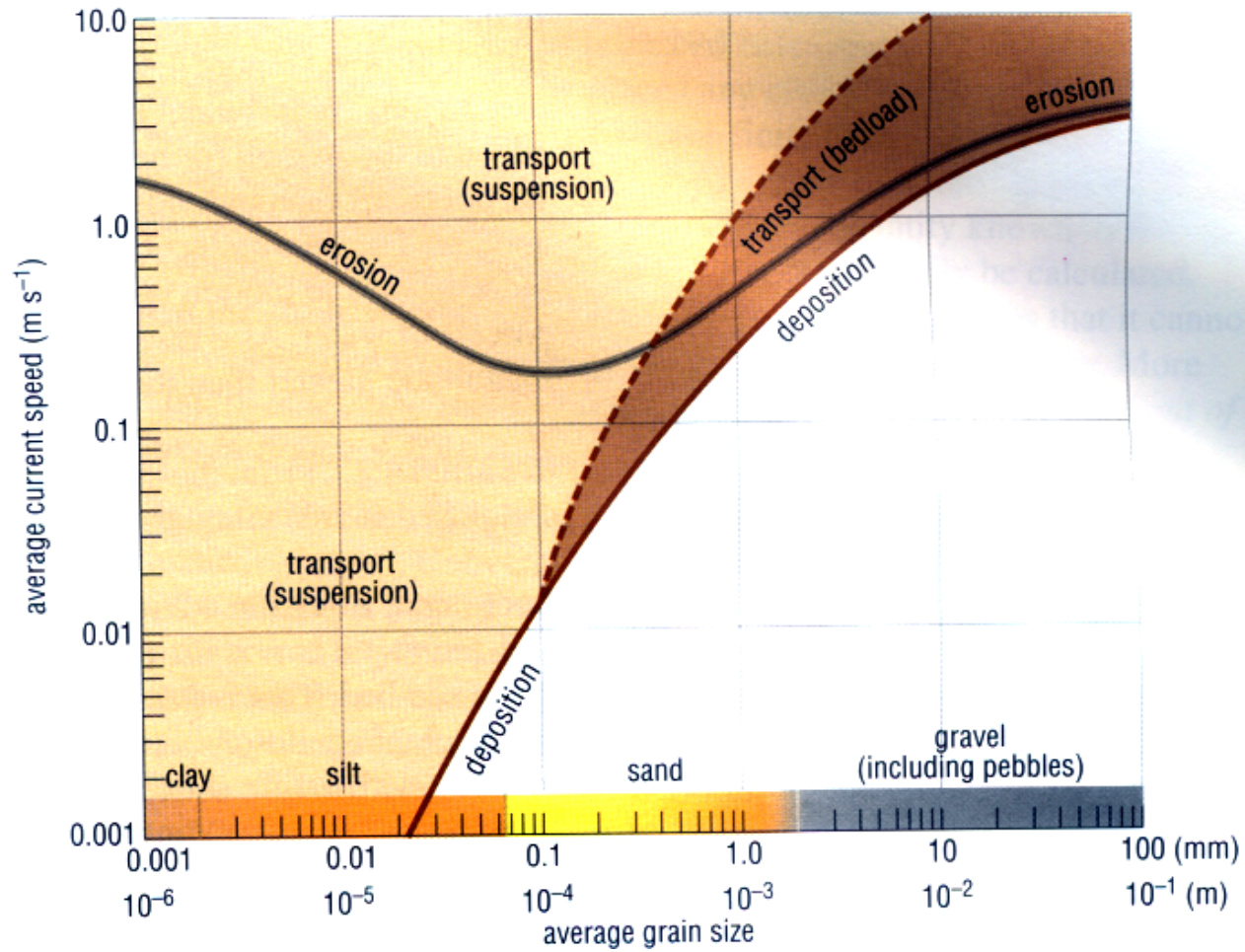
\bar{u}



(b)

Size of grain

\bar{u}

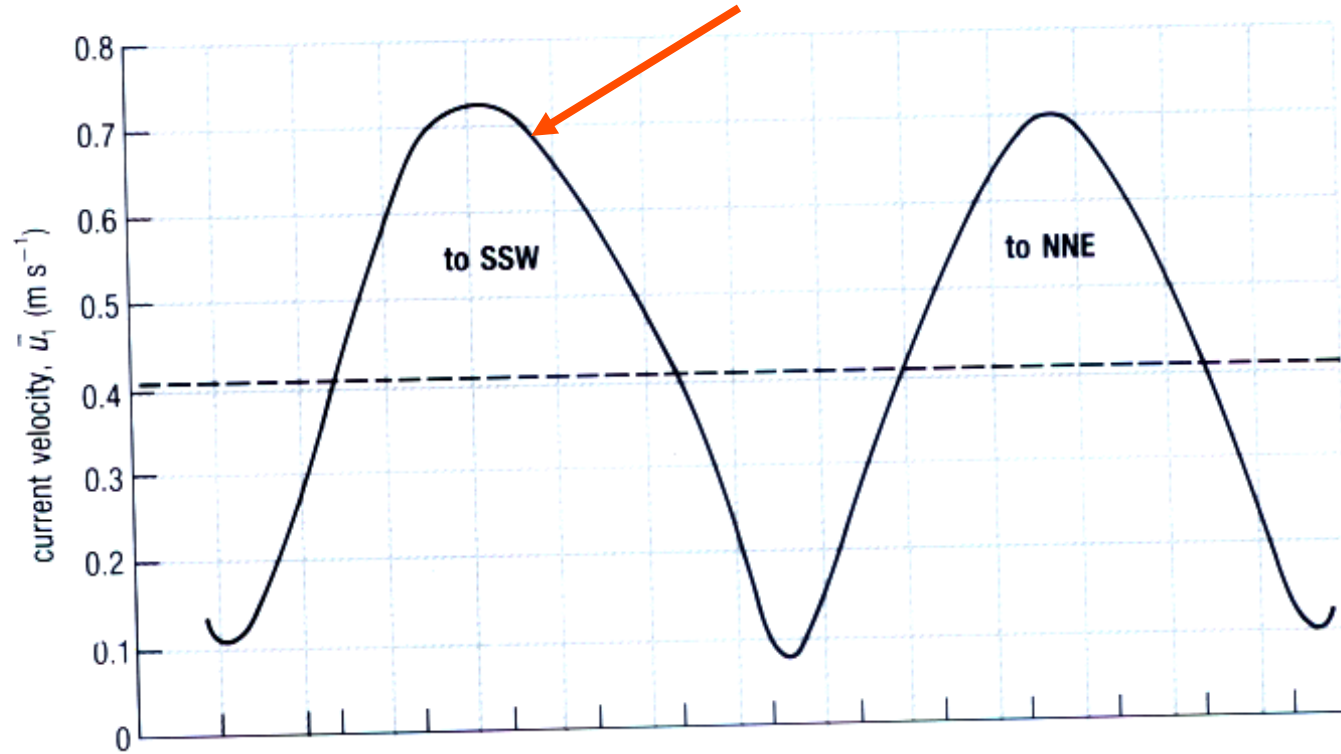


Size of grain

Tidal currents

(a)

Grain size > 0.3 mm moves



Rates of sediment transport

Sediment Flux →

A) Bedload

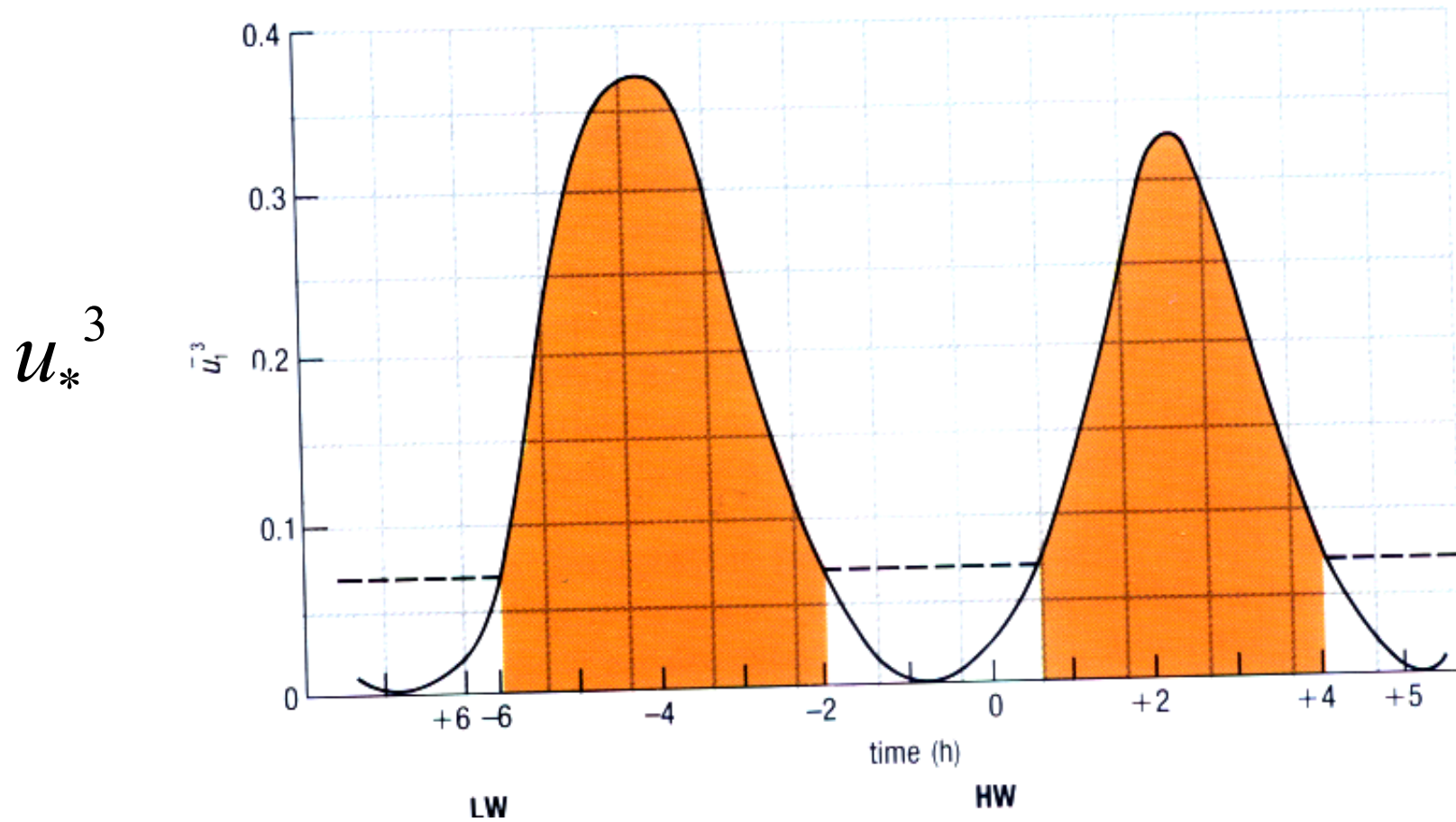
B) Suspended Load

Sediment Flux

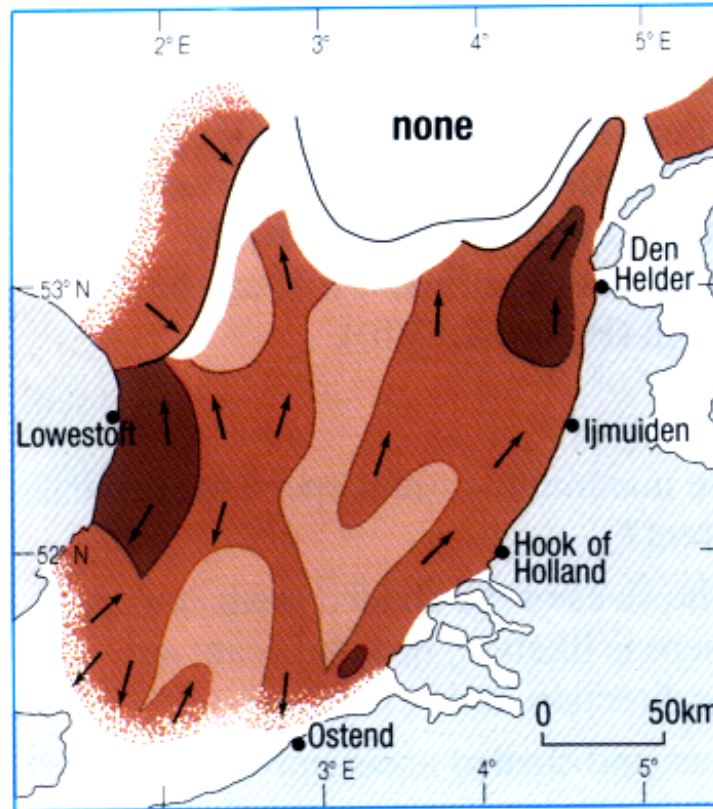
Bedload

$$q \propto \rho u_*^3$$

(b)



Maps of SAND TRANSPORT

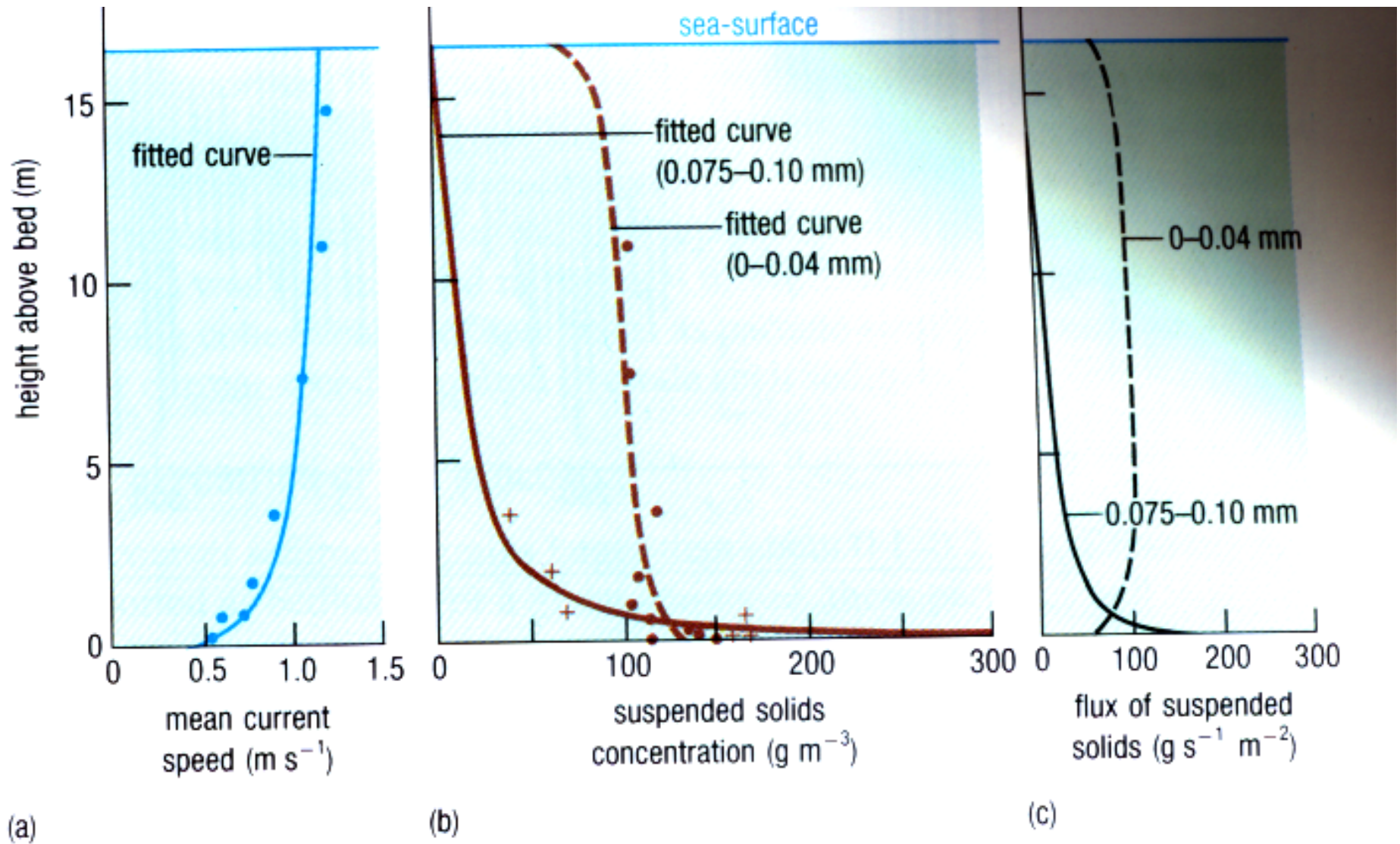


theoretical net sand transport



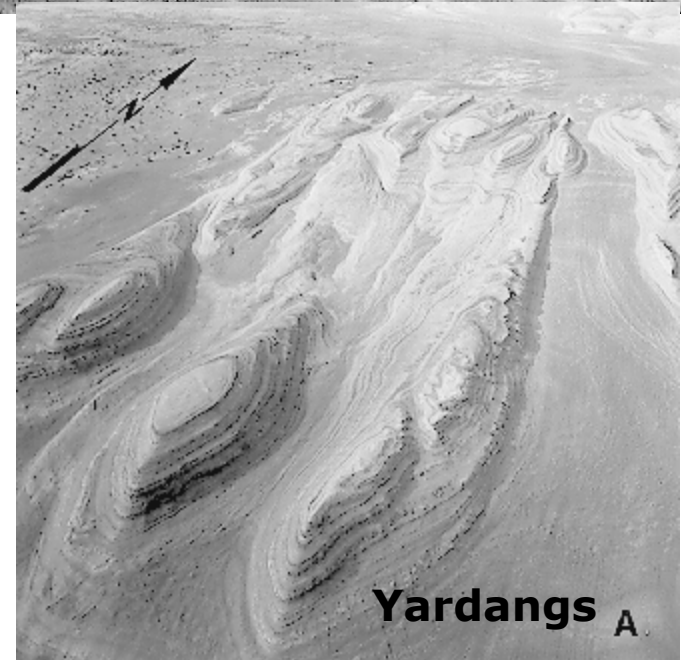
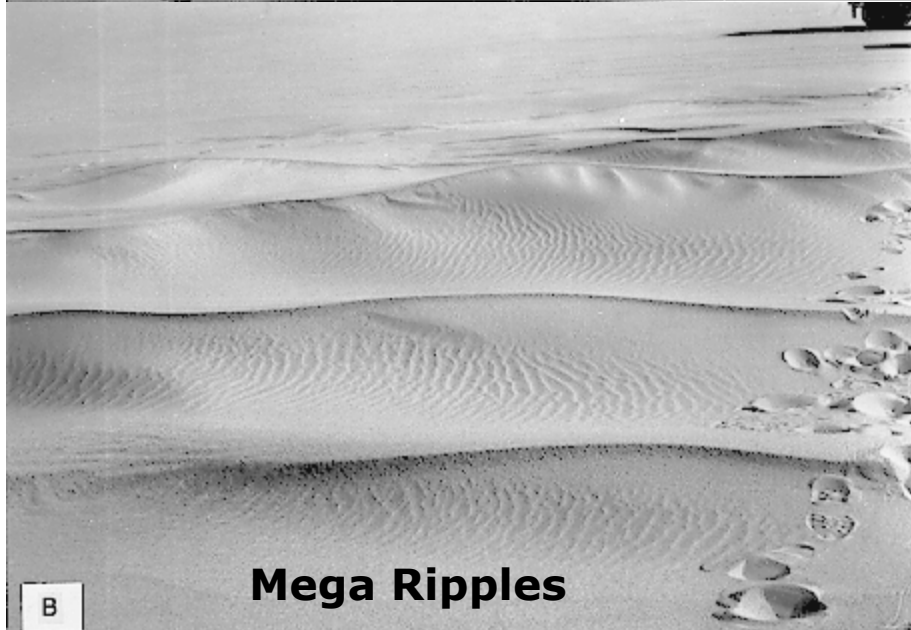
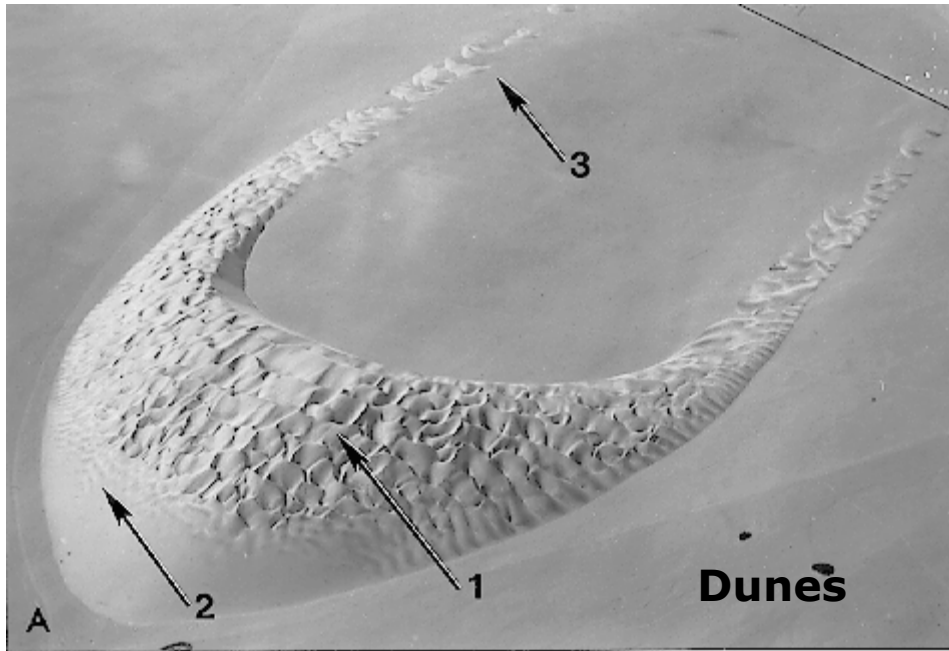
Sediment Flux

Suspended Load



Bed Forms

In the ocean the systematic study of bed forms started in 1940 with echo-sounding techniques



low



Ripples

slow current
0.6 mm grain size
need viscous sublayer

currents



height

wavelegnth



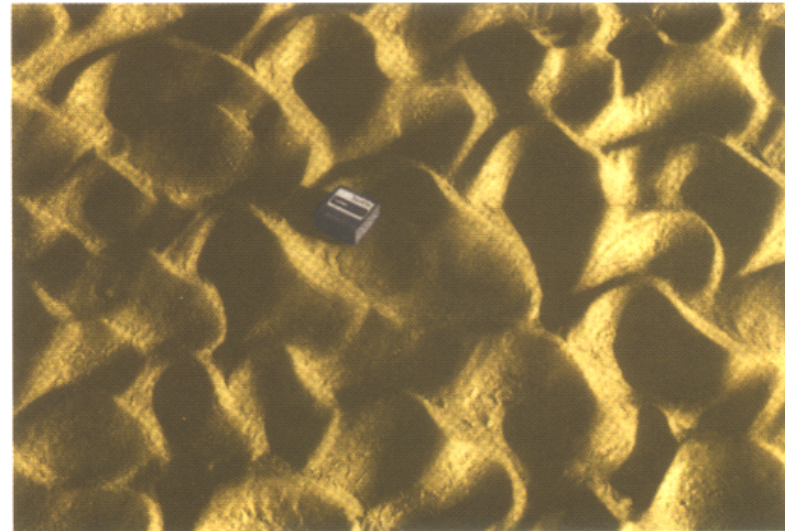
Mega Ripples

stronger currents
up to 1 meter high
disturbs ocean surface

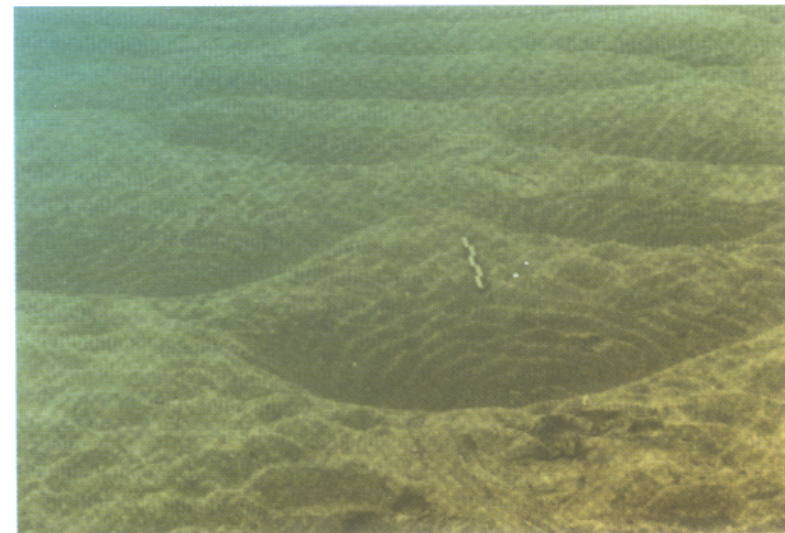
Sand Waves

stronger currents
up to 18 meter high
1 km wavelegnth

high

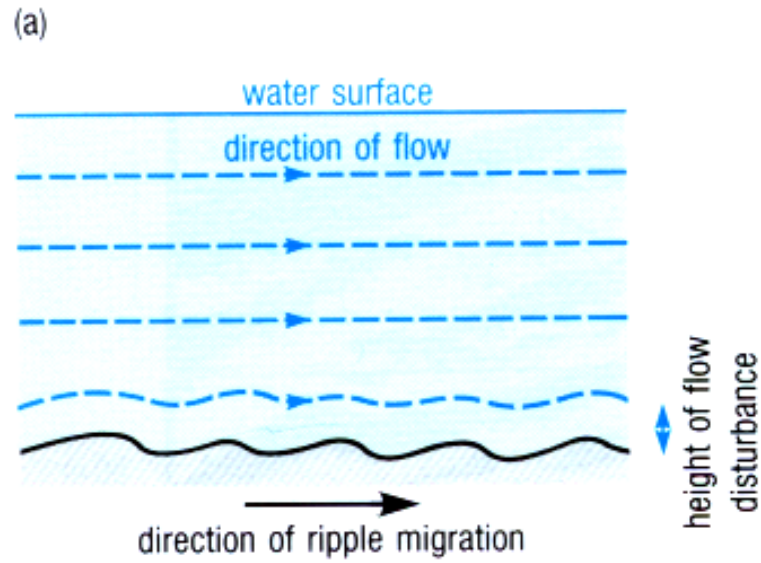


(a)

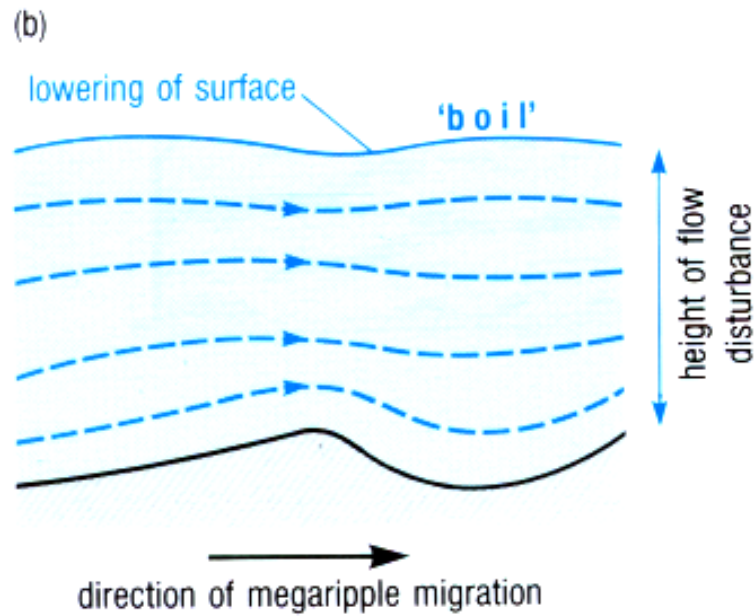


(b)

Ripples

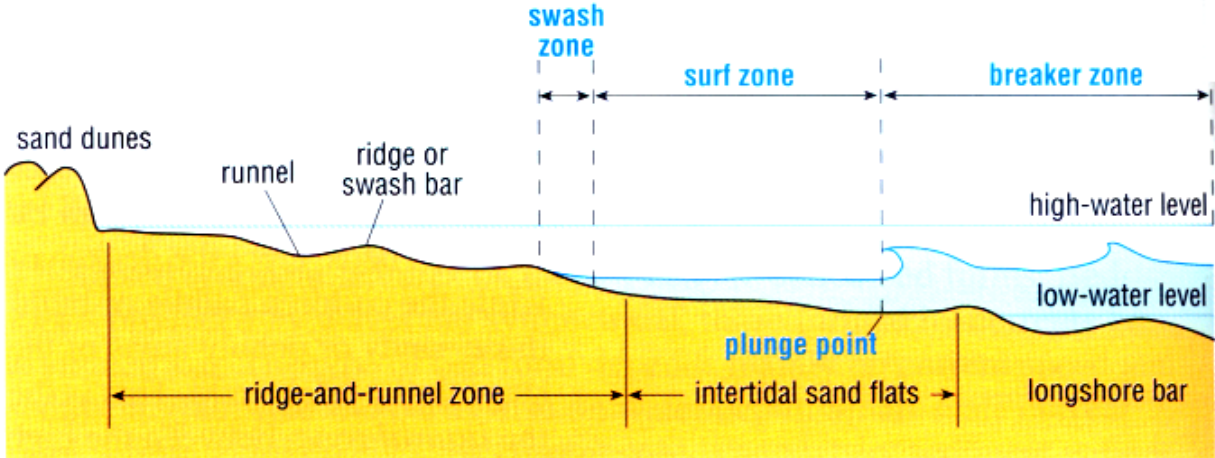


Mega Ripples



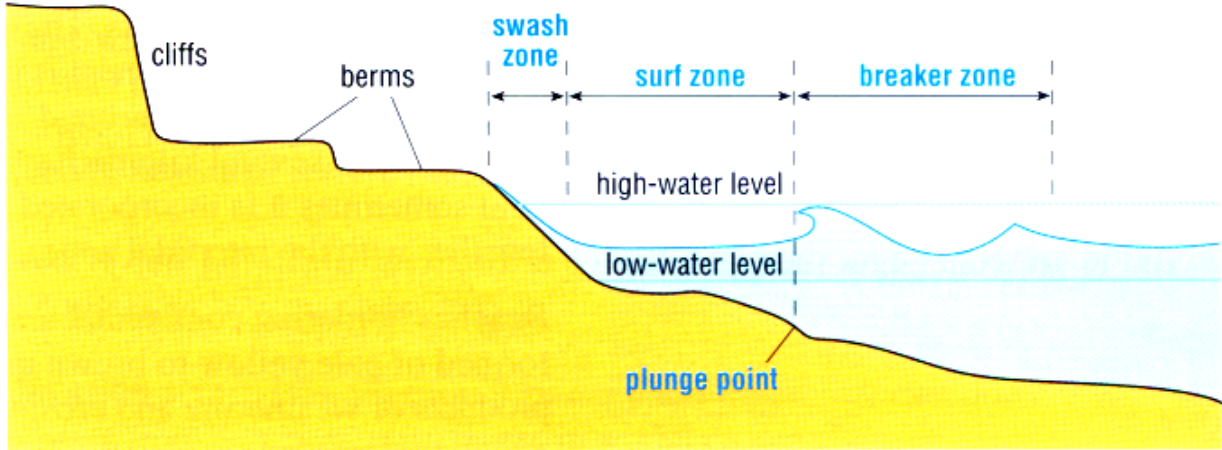
Beach Morphology

shallow slope



(a)

steep slope



(b)



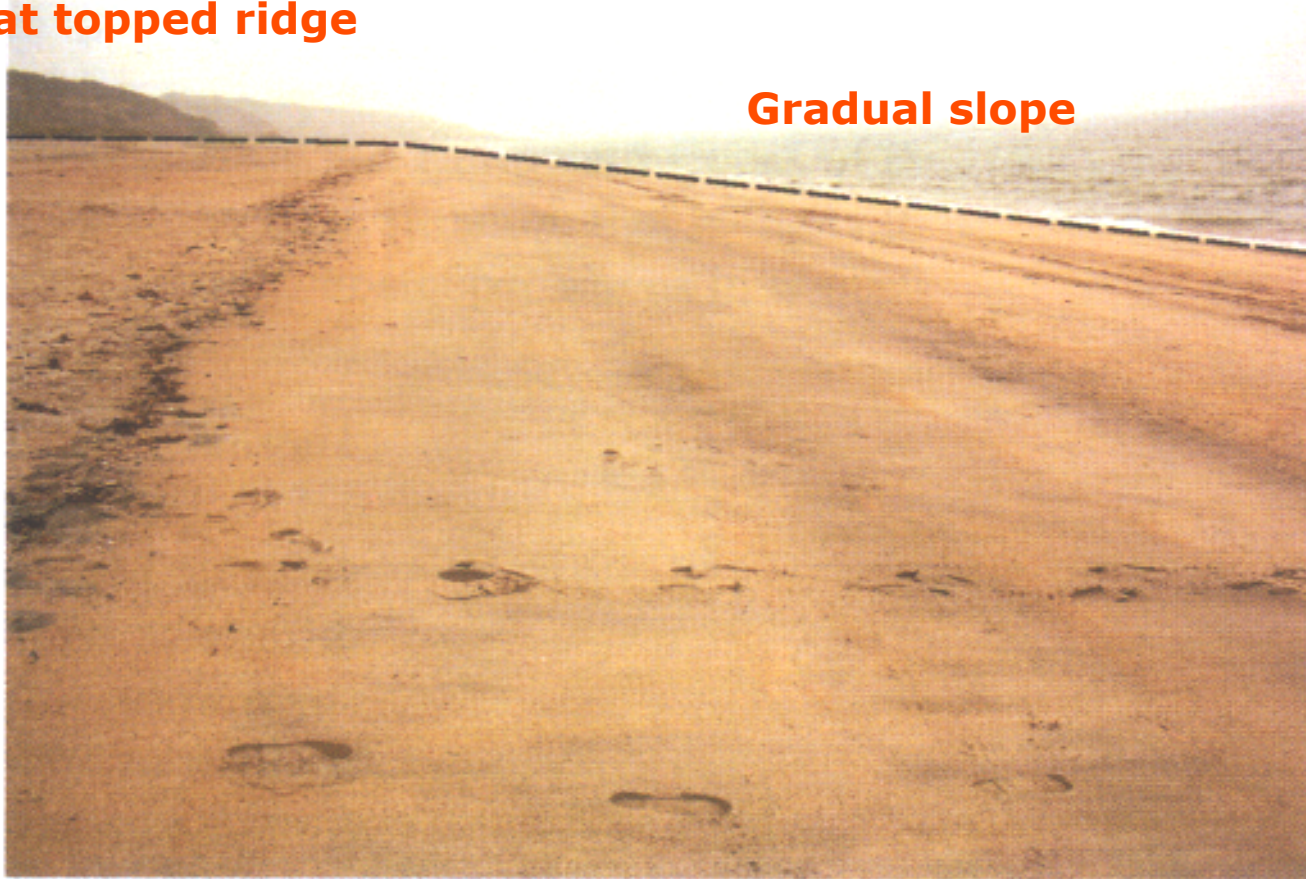
Surf Zone

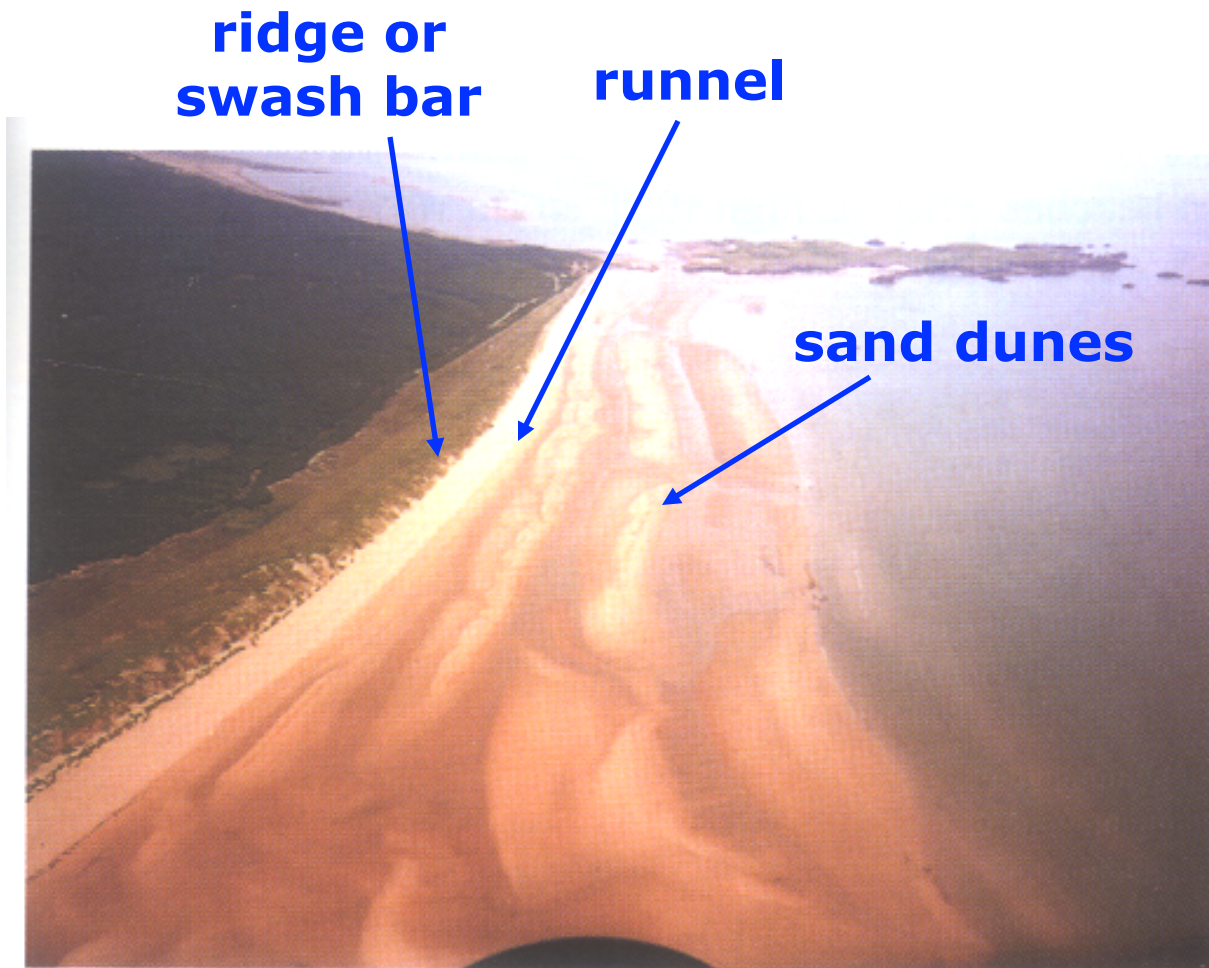
Breaker Zone

berm

Flat topped ridge

Gradual slope





**Gradual slope
intertidal sand flats**

**ridge or
swash bar**

runnel



(f)

Autumn



(a)

Winter



(b)

Winter



(c)

Summer



(d)